

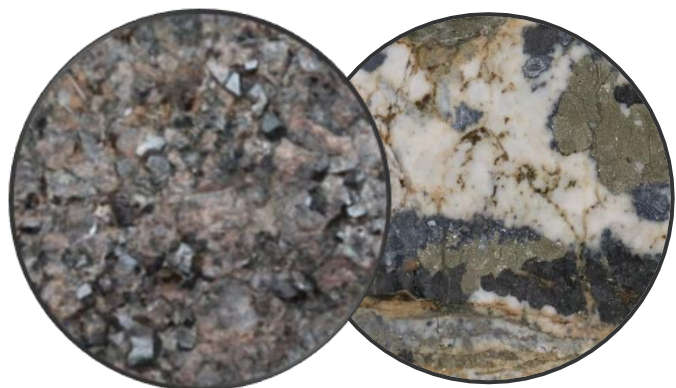
# Neues Potential

## First interim report

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## 1. Introduction

This document is the first internal report for the “Neues Potential” project in collaboration with the Sächsisches Landesamt für Umwelt, Landwirtschaft und Geologie (LfULG). This two-year project will develop a novel mineral systems model for the Eastern Erzgebirge (Germany) with a focus on magmatic-hydrothermal Li-Sn-W greisen, skarn and Ag-Pb-Zn vein-style mineralization related to late Variscan or post-orogenic granitoid intrusions. The results obtained will be used to constrain search space and to define exploration vectors for future mineral exploration in the Eastern Erzgebirge region.

In this internal report, we provide the status of the project and the progress of the six first months (from November 2022 to April 2023). The anticipated next steps for the upcoming six months are also presented.

The main tasks of the “Neues Potential” project include:

- acquisition and compilation of structural and geochemical data
- integration and synthesis of multiple approaches and datasets (field observations, petrography, structural analyses, fluid inclusion, geochronology)
- update on the geological overview of the magmatic and magmatic-hydrothermal systems, and essential knowledge gaps in the study area based on historic and new data
- building of a conceptual metallogenic model for the late-variscan magmatic and epithermal systems of the eastern Erzgebirge
- publication of new datasets as well as novel tectonic and metallogenic concepts.

These tasks are organized into six work packages that constitute the structure of this internal report:

- (1) Geochronology and geochemistry (work package 1): utilization of a wide range of analytical tools (e.g., microthermometry, LA-ICP-MS) to clarify the genetic relationship of diverse ore deposit types such as greisen and Ag-Pb-Zn epithermal veins. For this purpose, a wide selection of samples from historic geoscientific collections, exploration companies drill cores and surface outcrops are investigated in order to cover as much as possible the study area.
- (2) Structural and tectonic interpretation (work package 2): a structural geological and tectonic interpretation will be developed by synthesis of available datasets and

integration of new geochronology data to provide a conceptual structural framework for ore deposit genesis.

- (3) Development of an integrated mineral system (work packages 3 to 5): The development of one or more hypothesis(es) on the source, transport paths, and traps for ore deposits using the mineral systems framework.
- (4) Publication and dissemination (Work Package 6): results will be made available via the publication in “open access” peer-review journals and LfULG scientific publications.

## 2. WP 1 - Geochronology and geochemistry

Disentangling the sequence of mineralising events, and the spatio-temporal relationship between the magmatic and epithermal mineral systems, is key to the development of a comprehensive mineral system model for the Eastern Erzgebirge. This work package aims to review and collect new geochronology, mineral chemistry and fluid inclusion data to address these questions.

### 2.1. Current progress

#### *2.1.1. Data compilation and identified knowledge gaps*

Literature research was carried out to compile the geochronological data previously published in German and international scientific literature up to the present day (April 2023). Ages for the East Erzgebirge are compiled in the Appendix 1. Literature research illustrates an abundance of radiometric age data for many of the granitic and rhyolitic rocks of the Eastern Erzgebirge region, but also some very obvious gaps, for example, no radiometric age appears to be available for the Schellerhau granite in the central part of the Altenberg-Teplice caldera. Furthermore, there is only one rather poorly constrained age of  $278 \pm 20$  Ma (Rb-Sr on sphalerite) available for the Ag-Pb-Zn epithermal veins in the Freiberg district (Ostendorf et al., 2019). Although this age is deemed robust, its large error does not allow for confident correlation of vein formation with regional magmatic systems. Multiple ages are available for the Sn-Li-greisen deposits inside or in the vicinity of the Altenberg-Teplice caldera, but some of these are conflicting – and not in agreement with field geological evidence. The above illustrates that the acquisition of additional ages for the mineral deposits and magmatic rocks of the Eastern Erzgebirge is needed. In the focus of the present study is the age-dating of smaller greisen occurrences and veins that will help better constrain the timing of ore deposits in the study region.

### 2.1.2. Sample acquisition and preparation

A selection of ~50 samples were gathered from the TUBAF mineral collection, LfULG archives and recent Excellon drill cores. Of these, 32 were chosen to constitute a first batch of samples for geochronology (Tables 2-1 and 2-2). From the Ag-Pb-Zn epithermal veins of the Freiberg and Meißen districts, 16 samples hosting carbonates (e.g., dolomite-ankerite, Mn-bearing calcite, rhodochrosite) were selected. The localities sampled are Kleinvoigtsberg, Reinsberg, Brand-Erbisdorf, Bräunsdorf, Grauer Wolf, Zug, Reichenbach and Halsbrücke near Freiberg and Scharfenberg near Meißen (Fig. 2-1). These samples are mainly from previous Ph.D and M.Sc. projects, so their mineral paragenesis and association with mineralization is well documented. From Sn-greisen bodies and associated Sn-quartz veins, 16 samples were selected from historic mining districts of Zinnwald, Altenberg, Bärenfels, Niederpöbel, Lauenstein, Schmiedeberg, Panorama Höhe, Sachsenhöhe and Krupka. The sampling locations are illustrated in Figure 2-1.

Table 2-1. Cassiterite-bearing samples of Sn greisen and veins from the surrounding of the Altenberg-Teplice caldera and cassiterite-bearing samples of the Ag-Pb-Zn epithermal veins of the Freiberg district.

Sample ID	Depth	District	Subdistrict	Mineralogy	Host-rock	Collection
MG_Znn_66533		Zinnwald	Zinnwald	Cassiterite, wolframite	Altered granite	TUBAF
65775		Altenberg	Altenberg	Coarse tin concentrate	Altered granite	TUBAF
MG_Znn_43652		Zinnwald	Zinnwald	Cassiterite, zinnwaldite, quartz	Altered granite	TUBAF
MG_Znn_43672		Zinnwald	Zinnwald	Cassiterite, zinnwaldite	Altered granite	TUBAF
MG_Znn_43697		Zinnwald	Zinnwald	Cassiterite, wolframite	Altered granite	TUBAF
MG_Znn_43659		Zinnwald	Zinnwald	Cassiterite (massive, smooth surfaces), zinnwaldite, quartz	Altered granite	TUBAF
CRFG4		Freiberg	Kleinvoigtsberg*	Pyrite, proustite, galena, quartz, cassiterite		TUBAF
MG_HB_51003		Freiberg	Halsbrücke	Cassiterite (?), galena		TUBAF
MG_BA_66818			Bärenfels	Cassiterite		TUBAF
43245				Wolframite, cassiterite	Altered granite	TUBAF
MG_SK_SS18875			Lauenstein	Cassiterite		LfULG
MG_SC_SS4024			Schmiedeberg	Cassiterite in quartz vein	Altered granite	LfULG
MG_KP_RS8653		Krupka	Krupka	Cassiterite		LfULG
MG_NDP_66044			Niederpöbel	Cassiterite, quartz		TUBAF
14-23b			Sadisdorf	wolframite		TUBAF



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65803

Altenberg

Altenberg\*\*

Cassiterite

TUBAF

\* Alte Hoffnung Gottes mine

\*\* Zinnklüfter vein, „Zwitterstocks tiefen Erbstolln“

Table 2-2. Carbonate-bearing samples from Ag-Pb-Zn epithermal veins of the Freiberg and Meißen districts.

Sample ID	District	Subdistrict	Description	Mineralogy	Host-rock	Collection
52924A	Freiberg	Reinsberg	Emanuel Erbstolln, Reinsberger Glück Morgengang, über 4. Gezeugstrecke in Nordost	Quartz, carbonate, proustite/pyrite, pyrite		TUBAF
LSFG024	Freiberg	Kleinvoigtsberg	Alte Hoffnung Gottes, Heinrich Stehender	Ag-Sb-sulfides, quartz		TUBAF
51538b	Freiberg	Brand-Erbisdorf	Himmelsfürst Fundgrube, Jupiter Stehender, 3. Gezeugstrecke, aus 1. Benjaminschacht in Süd	Galena, sphalerite, pyrite, ankerite, quartz	Gneiss	TUBAF
MG_BRD_01_353.30_1	Freiberg	Bräunsdorf		Quartz, carbonate, arsenopyrite, pyrite		EXCELLON
MG_GWO_030_272.65_1	Freiberg	Grauer Wolf		Quartz, carbonate, arsenopyrite, pyrite		EXCELLON
MG_REI_027_179.10_1	Freiberg	Reichenbach		Quartz, carbonate, arsenopyrite, pyrite		EXCELLON
MG_GWO_038_78.66_1	Freiberg	Grauer Wolf		Quartz, carbonate, arsenopyrite, pyrite		EXCELLON
MG_GWO_038_61.15_1	Freiberg	Grauer Wolf		Calcite (normal fault infill)		EXCELLON
MG_GWO_038_48.0_1				Quartz, carbonate, arsenopyrite, pyrite		
MG_GWO_040_392.8_1						
MG_KVB_52712	Freiberg	Kleinvoigtsberg	Alte Hoffnung Gottes Erbstolln, Frisch Glück Stehender, 32 Lachter vom Gottliebschacht in Süd	Galena, pyrite, carbonate		TUBAF
MG_SFB_53248	Meissen	Scharfenberg				TUBAF
52936B	Freiberg	Reinsberg	Emanuel Erbstolln, Reinsberger Glück	Quartz, carbonate,		TUBAF

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			Morgengang, hartes Trum	chalcopyrite, acanthite	
MG_Z_WSt1	Freiberg	Zug	Wilhelm Stehender (Baustelle)	Sulfides, quartz, carbonate	Jens Gutzmer private collection
MG_BE_50881	Freiberg	Brand-Erbisdorf	Habacht Fundgrube, Hangendes Trum des Ludwig Stehenden bei 1. Gezeugstrecke, 90 Lachter vom Jung Schwarzfärber Spat in Nord	Boulangerite, galena, sphalerite, owyheeite, rhodocrosite, quartz	TUBAF
MG_BE_51145	Freiberg	Brand-Erbisdorf	Vergnügte Anweisung Fundgrube samt Reussen Fundgrube, Friedrich Stehender (carbonate breccia)	Galena, ankerite, calcite, pyrite, quartz	Gneiss TUBAF

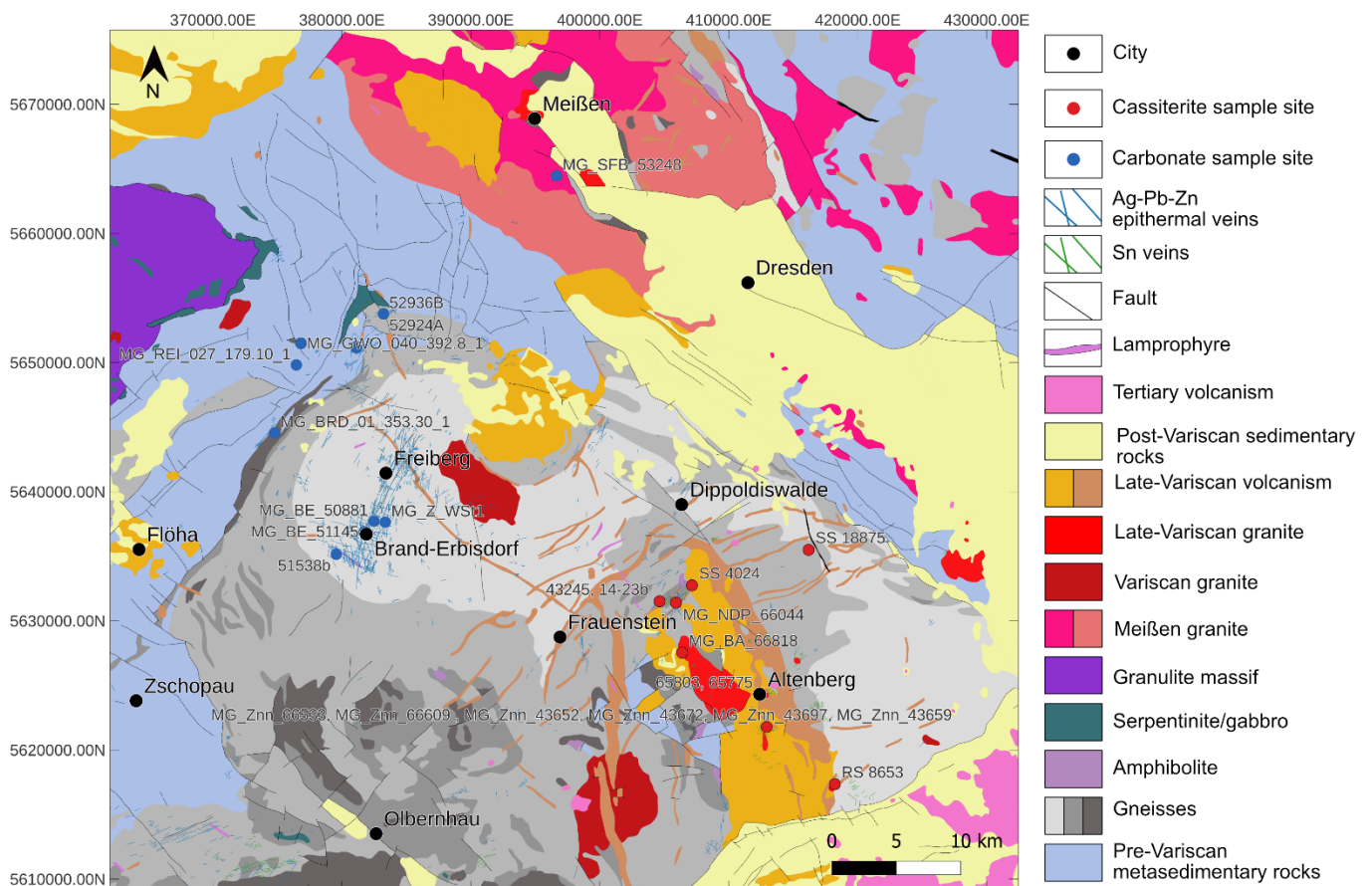


Figure 2-1. Sampling strategy for carbonates (in blue) and cassiterites (in red) from Ag-Pb-Zn epithermal vein and greisen bodies and veins. Simplified geological map modified after Hoth et al. (1980).

### *2.1.3. Geochronological measurements*

The first batch of 32 samples was analysed on the 15 and 16 of December 2022 at the FIERCE laboratory at the Goethe University in Frankfurt (Germany). The U-Pb datation by La-ICP-MS is a method that has been successfully developed and used over the last decade on a large variety of minerals (e.g., garnet; Burisch et al., 2019a; Reinhardt et al., 2022). The specificity of this method is that it can be applied to low-U mineral phases such as carbonate (Burisch et al., 2017; Roberts et al., 2020; Guilcher et al., 2021). Results are expected within the next few months and will be reported in the next interim report.

### *2.1.4. Development of integrated time-space tectonic diagram*

The ages compiled from published literature for magmatism, volcanism and ore deposits in the eastern part of the Erzgebirge are presented in Figure 2-2. This will be progressively updated over the course of the project, as new data become available.

## 2.2. Current challenges

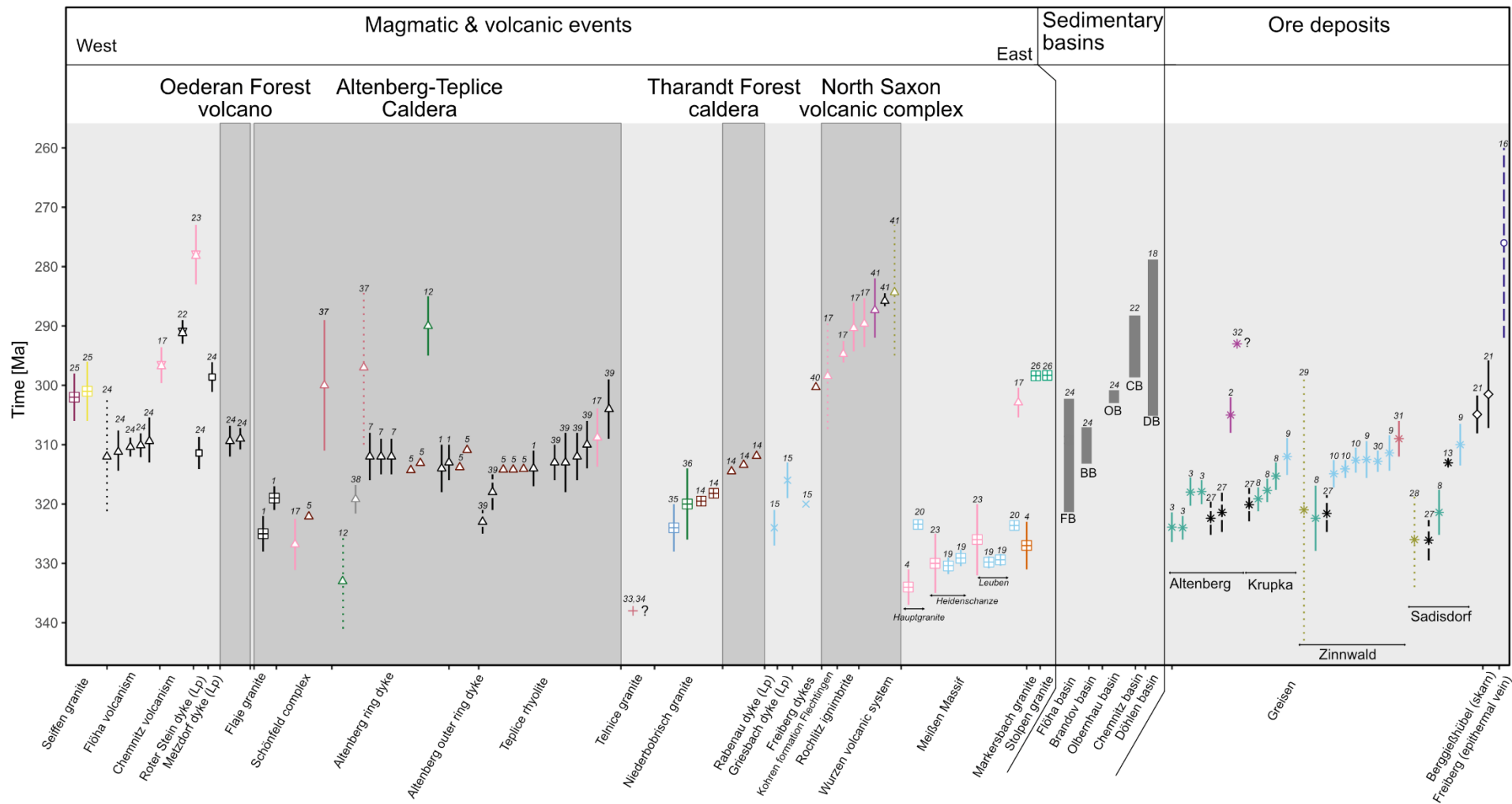
Several difficulties were encountered during the initial stages of this study. These are reported below for reference.

1. There is an obvious lack of suitable sample material in accessible collections and archives for many of the smaller vein deposits in the Eastern Erzgebirge. Minor Ag-Pb-Zn veins are reported by Baumann (1965) near Sayda, Reichenau, Frauenstein, Schönfeld, Naundorf, Dippoldiswalde and Edle Krone (Fig. 2-3). Yet, no sample material from these sites is available. Therefore, no samples could be included in the first batch of samples for geochronology. Sampling from dumps, outcrops or visitor mines will be carried out in the coming months for geochronology and fluid inclusion study purposes.

2. For the greisen deposits, samples from the LfULG hand specimens collection for the localities of Bärenstein, Glashütte, Lauenstein, Schönfeld and Dippoldiswalde mostly consist of non-mineralized greisen samples. Samples from these localities in the TU Bergakademie Freiberg geoscientific collection are scarce. Additionally, no greisen samples were collected from the Falkenhain district because access to potential sample material was not provided by Deutsche Lithium GmbH.

3. An extensive search was conducted for cassiterite-bearing samples from the Pb-Zn-Ag epithermal veins in the Freiberg district to complement our attempt to date carbonate veins (as carbonates often do not give robust ages). Baumann (1965) reported euhedral cassiterite crystal in several samples from thin sections that are now stored in the ore deposit collection

of the TU Bergakademie Freiberg. However, up to today, no access to these thin sections has been provided by Prof. Thomas Seifert (currently responsible for the collections). Instead, we are actively reaching out to museums and private collectors to find suitable alternative cassiterite samples.



Event	Method (ages)	Abbreviation	Method
⊠ Basin	— Robust	OB Olbernhau basin	● Ar-Ar (biotite, amphibole, Li-mica, muscovite)
□ Dyke	- - - Questionable	FB Flöha basin	● K-Ar (amphibole, biotite, Li-mica)
○ Epithermal vein	⋯ Passable	BB Brandov basin	● K-Ar (siderite)
* Greisen		DB Döhlen basin	● Pb-Pb evaporation zircon
× Lamprophyre		CB Chemnitz basin	● Rb-Sr isotopes sphalerite
⊞ Pluton			● Rb-Sr whole-rock
◇ Skarn	<b>Literature</b>		● Re-Os molybdenite
△ Volcanism	1 reference		● Sm-Nd (fluorite, wolframite)
			● U-Pb LA-ICP-MS (cassiterite, garnet, zircon)
			● U-Pb LA-ICP-MS (monazite, xenotime, uraninite)
			● Th-U-Pb (monazite)
			● U-Pb CA-ID-TIMS zircon
			● U-Pb SHRIMP zircon
			● U-Th-Pb LA-ICP-MS zircon
			● U-Pb evaporation zircon

Figure 2-2. Time-Space diagram compiling all ages available classified by the dating method used. References: 1. Tomek et al. (2021), 2. Gerstenberger et al. (1989), 3. Romer et al. (2007), 4. Hofmann et al. (2009), 5. Tichomirowa et al. (2022), 6. Opluštil et al. (2016), 7. Tomek et al. (2019), 8. Akerman et al. (2017), 9. Seifert et al. (2016), 10. Seifert et al. (2011), 11. Atanasova et al. (2012), 12. Kempe et al. (1999), 13. Leopardi et al. (*submitted*), 14. Breitzkreuz et al. (2021), 15. Von Seckendorff et al. (2004), 16. Ostendorf et al. (2019), 17. Hoffmann et al. (2013), 18. Zieger et al. (2019), 19. Wenzel et al. (1997), 20. Sharp et al. (1997) in Müller (2011), 21. Burisch et al. (2019a), 22. Luthardt et al. (2018), 23. Nasdala et al. (1999), 24. Lócse et al. (2019), 25. Förster and Rhede (2006), 26. Käßner et al. (2021), 27. Zhang et al. (2017), 28. Kempe and Belyatsky (1997), 29. Höndorf et al. (1994), 30. Neßler et al. (2016), 31. Dolejš and Štemprok (2001), 32. Haack (1990), 33. Štemprok et al. (2003), 34. Klomínský et al. (2010), 35. Förster et al. (1998), 36. Tichomirowa (1997), 37. Müller et al. (2005), 38. Romer et al. (2010), 39. Casas-Garcia et al. (2019), 40. Marion Tichomirowa (personal communication), 41. Wendt et al. (1995)

### 2.3. Next steps

The goals for the next six months are as follows:

A second batch of cassiterite from the Ag-Pb-Zn veins is being prepared to complement the carbonate dating of December 2022. The new batch of samples will be dated around August 2023 using the same method (U-Pb LA-ICP-MS).

Field work is scheduled to sample Ag-Pb-Zn veins located in the surrounding of the Altenberg-Teplice and the Tharandt caldera. New sampling is necessary to tackle the questions of fluid nature, paleodepth, and geochemistry of epithermal veins from other localities than Freiberg (e.g., Frauenstein, Edle Krone and other localities mentioned in section 2.2, Fig. 2-3). These areas have been identified as key knowledge gaps that potentially demarcate the boundary between the Freiberg and Altenberg-Teplice mineral systems.

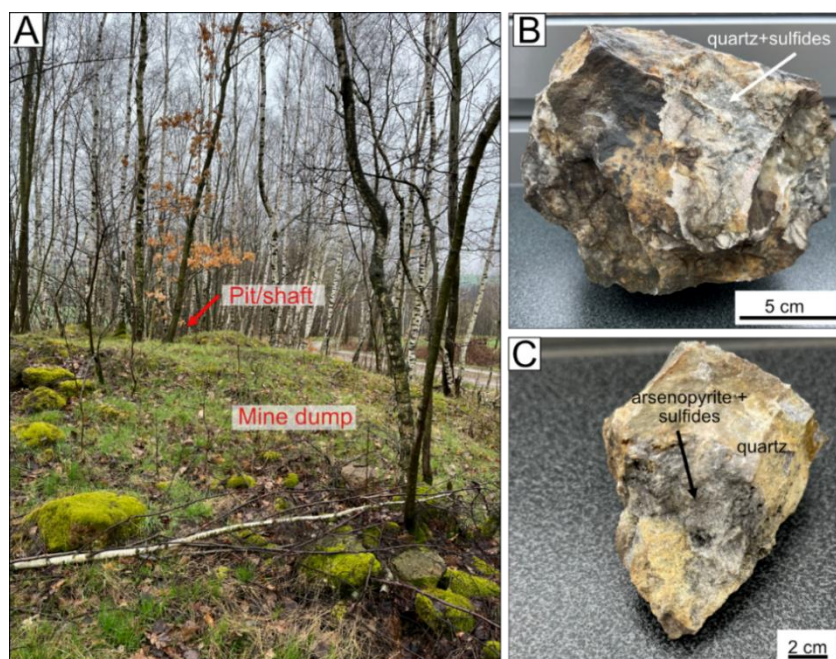


Figure 2-3. (A). Mine dump next to an abandoned pit/shaft in Mittelreichstädt (Fig. 2-4, near Reichstädt, Dippoldswalde district, coordinates UTM zone 33: 403482.4, 5634737.0) located on a Ag-Pb-Zn epithermal veins occurrence. (B-C). Samples collected from the mine dump.

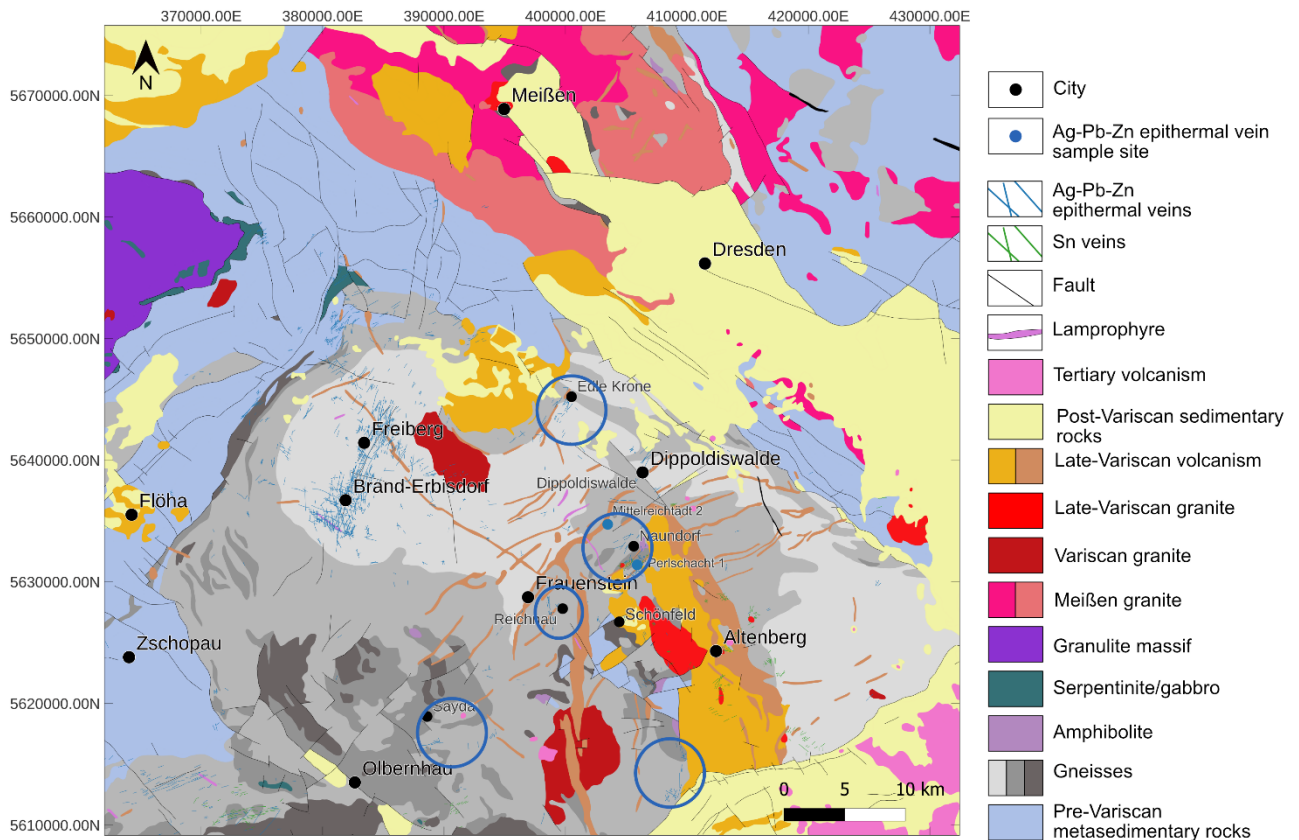


Figure 2-4. Sampling strategy for Ag-Pb-Zn epithermal veins for future geochronology and geochemistry/fluid inclusion analysis. Sampling is focused on districts in the surroundings of the Altenberg-Teplice caldera, Dippoldiswalde district and Tharandt caldera. Simplified geological map modified after Hoth et al. (1980).



### 3. WP 2 - Structure and Tectonics

Crustal and local-scale structures, and their interaction with the broader tectonic setting, are a key control on the localization of magmatic-hydrothermal ore deposits. This work package focuses on (1) reviewing the range of possible tectonic settings for the Eastern Erzgebirge mineral system, especially after the peak Variscan metamorphism, and (2) using mapped and inferred structures to provide geometric constraints on flow pathways that focused mineralizing magmas and fluids during post-Variscan mineralizing events. Potential links between magmatic and epithermal mineralization and the Tharandt and Altenberg-Teplice caldera systems will also be considered.

#### 3.1. Current progress

##### *3.1.1. Literature review and tectonic context*

The Eastern Erzgebirge (EE) is a part of the Saxothuringian unit (STU), which outcrops along the Variscan belt and forms the northern margin of the Bohemian Massif. Over the last decades, diverse concepts for the evolution of the Saxothuringian unit have been proposed. The structural complexity of the whole Bohemian Massif has led many authors to the conclusion that it resulted from several accretionary collisions between various microplates/terranes, however interpretations of the nature and timing of these collisions differ. These terranes consist of Upper Proterozoic to Upper Devonian rocks, and their geochemistry indicates different origins, including ocean floor (ophiolite), island arc, shelf, and active continental margin (e.g., Oliver et al., 1993). The first amalgamation has been suggested to have occurred along the Tornquist line, during the upper Ordovician time (~460–445 Ma), as some relics of Caledonian orogeny have been found in Poland (Oliver et al., 1993; Cymerman, 2000). The second suggested amalgamation coincides with Variscan orogenesis at Middle to Upper Devonian time (~390–360 Ma).



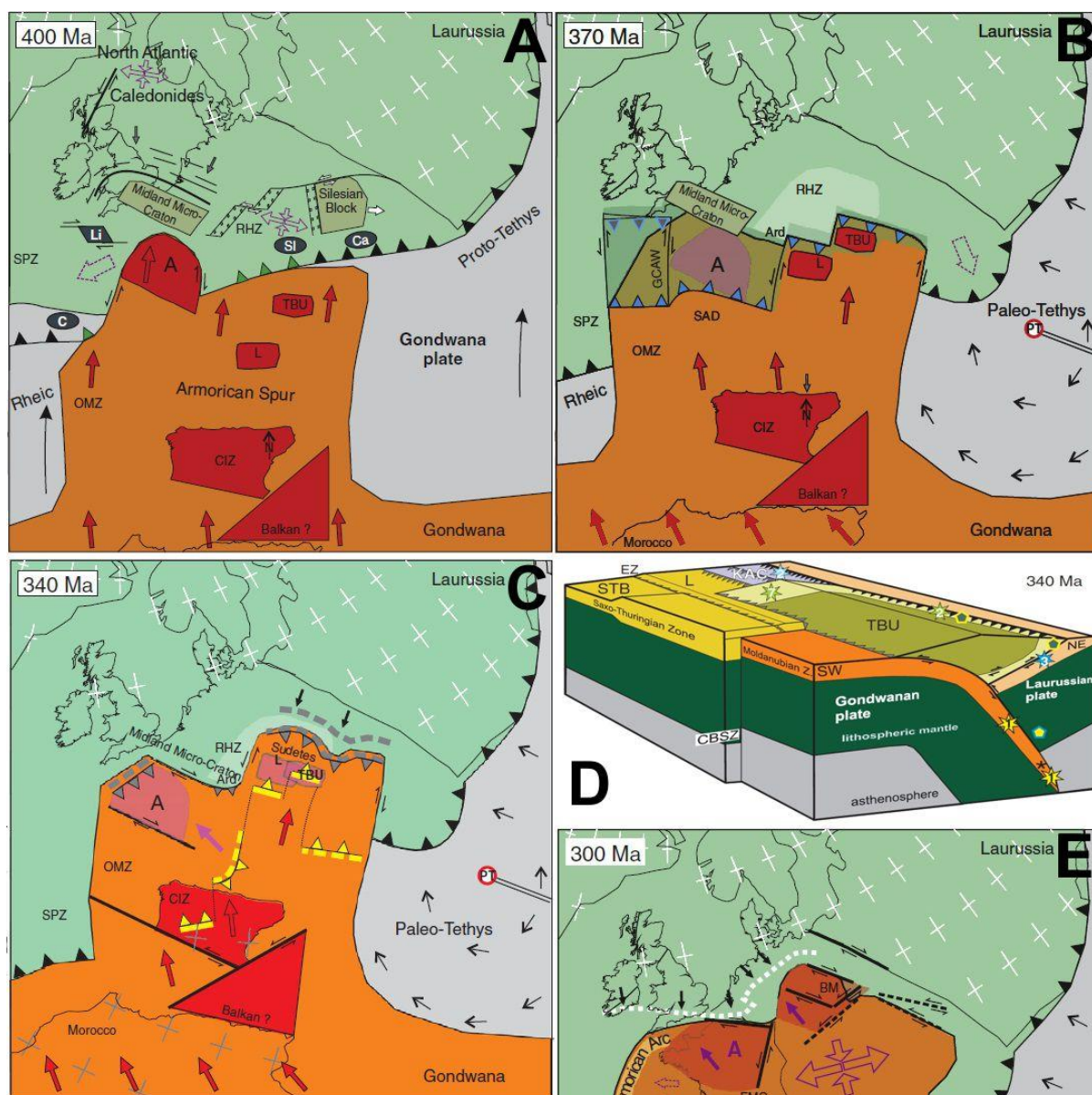


Figure 3-1. Evolution of Variscan units/terrane proposed by Kröner & Romer (2013): Li, C, Si, Ca = Recognized ophiolites in the upper (Laurussian) plate representing trans-tensional and back-arc basins at around 400 Ma, A = Armorica, TBU = Teplá-Barrandian unit, L = Lusatia, CIZ = Central Iberian Zone, OMZ = Ossa Morena Zone, RHZ = Rheno-Hercynian Zone, GCAW = Galicia-Cornwall Accretionary Wedge, Ard = Ardennes, SPZ = South Portuguese Zone, FMC = French Massif Central, BM = Bohemian Massif, PT = Euler Pole in the Paleotethys Ocean.

Kröner & Romer (2013) suggest that three separate subduction systems were developed during the accretion of smaller (Cadomian) terranes belonging to the Variscan orogeny: (1) Subduction zone I (410–380 Ma) with main subduction system dipping towards the north and being responsible for the initial accretion of the Armorican terrane with Laurussia (Fig. 3-1A). (2) Subduction zone II (380–360 Ma) when the main subduction jumps outwards to subduct southwards underneath the Teplá Barandien Unit (TBU; Fig. 3-1B). (3) Final subduction zone III (360–340 Ma) when another subduction developed on the SW margin of

the TBU dipping towards the NE (Fig. 3-1C, D). This subduction zone was situated SW from the TBU with a transition to collision with the Moldanubian unit (MDU) at around 340 Ma, which is also the age of peak metamorphism in the SU (Kröner et al., 2007). After that, the post-orogenic development was related mainly to a dextral shearing (Fig. 3-1E).

This model, however, does not explain: (1) why sedimentary units since Upper Famennian (360 Ma) are horizontal and lie discordantly on the older Devonian sequence, as structurally and paleontologically documented in the Nepasice borehole near Hradec Králové (Čech et al., 1989; Chlupáč & Zikmundová, 1976), which implies that collision was already over at that time; (2) Subduction system III does not explain well the origin of the Central Bohemian Plutonic Complex, which was emplaced during ~354–337 Ma and has distinctive continental arc geochemistry (early calc-alkaline plutons and late ultrapotassic series; Žák et al., 2005), as the distance of plutons from the interpreted subduction zone trench (Fig. 3-2D) would extend from 20-100 km, which would require an extremely steep dipping slab or unreasonable amount of tectonic shortening. Moreover, the prolongation of the pluton is perpendicular to the MDU-TBU boundary.

Unlike Kröner & Romer (2013), Žák et al. (2014) consider that collision between STU and TBU occurred earlier, between ~380–346 Ma, along a south-dipping subduction zone (i.e., STU subducted beneath TBU at its NW margin) responsible for calc-alkaline plutons in the TBU and high-pressure metamorphism in the Erzgebirge (STU). Then, another WNW dipping subduction was responsible for the emplacements of ultrapotassic plutons in Bohemian Massif at around ~346–335 Ma. All this was followed by an overall dextral-tectonics regime between 335–315 Ma (Fig. 3-2).

Furthermore, none of the above models sufficiently explain the apparent oroclinal termination of the Variscan, with the STU apparently wrapping around the TBU and MDU in the east (Franke and Zelazniewicz, 2000). The rapid transition to post 340 Ma extension and subsequent dextral strike-slip on NW trending structures (e.g., the Elbe shear zone) are also poorly explained (e.g., Franke et al., 2014).

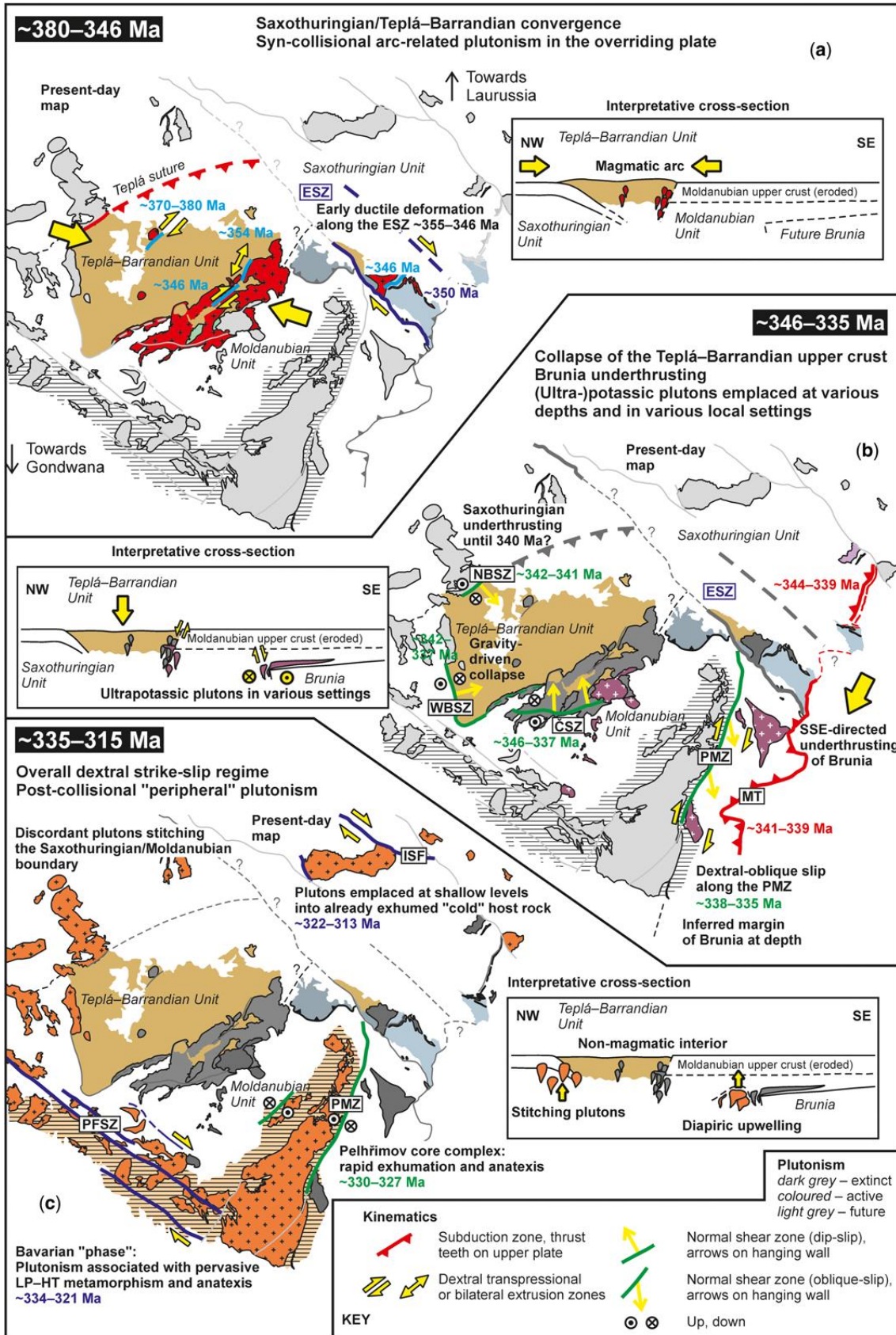


Figure 3-2. Evolution of Variscan plutons and subduction zones proposed by Žák et al. (2014).



Schulmann et al. (2014) suggest subduction of the TBU beneath STU quickly followed by a collision and exhumation of the lower crust. Unlike Kröner, he explains the Saxothuringian unit presence SE from TBU (See yellow and green areas representing STU in Fig. 3-3A) as exhumated lower crust (Figs. 3-3B-D). However, STU is affected by a lower grade of regional metamorphism than the MDU, which contradicts its interpreted lower-crustal source. Field observations also show that the STU lies structurally above the MDU, which we suggest makes their model is physically unrealistic as: (1) The suggested plastic behavior of the SU would require significant metamorphism to penetrate through the hard solid core of the MDU. However, STU outcropping inside MDU consist of shists, phyllites, marbles, cherts, quartzites, and lenses of amphibolites (i.e., low to medium grade metamorphism); (2) The STU is in general denser than the MDU, precluding standard models of lower crustal exhumation through buoyant doming.

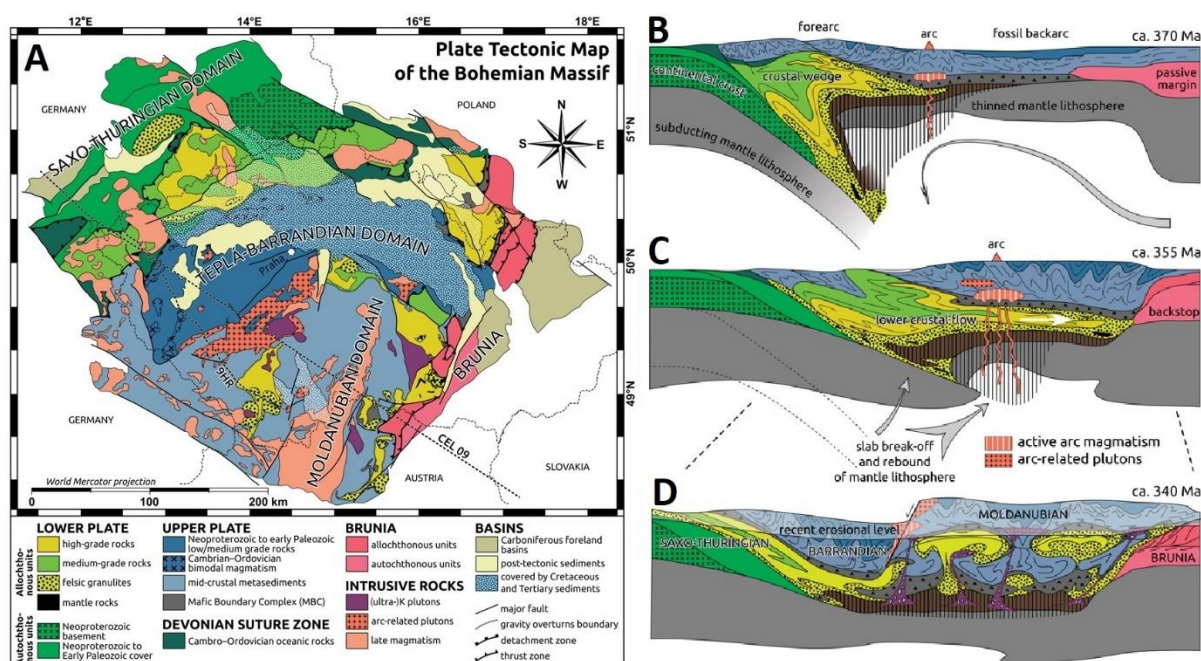


Figure 3-3. Evolution of Variscan orogen and subduction zone proposed by Schulmann et al. (2014).

Mazur et al. (2020) suggested that the main subduction zone lies far to the north, buried beneath the sedimentary cover in NE Germany between Avalonia terrain and Composite of Armorican Terranes, with a major post-collisional deformation from dextral strike-slips due to oroclinal folding of the eastern termination of the European Variscides (Fig. 3-4A). Unfortunately, their models do not give a mechanism for this oroclinal bending, nor do they precisely state the timing and, thus, does not relate the main calc-alkaline magmatic activity in Bohemian Massive (BM) to the Variscan subduction zone. However, like the other authors, they note the importance of dextral tectonic in the BM.

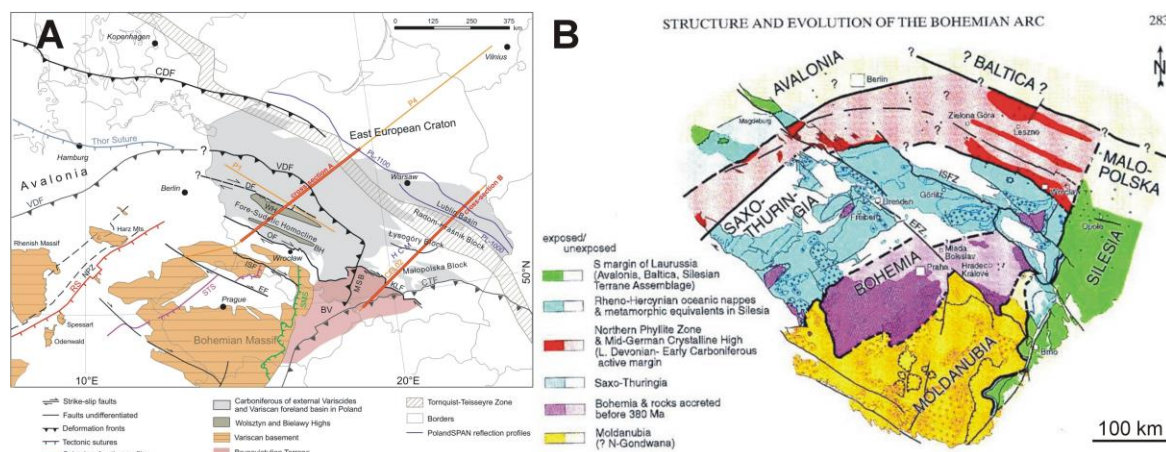


Figure 3-4. (A) Evolution of Variscan orogen and final deformation zone geometry proposed by Mazur et al. (2020): CDF = Caledonian deformation front, VDF = Variscan deformation front, CTF = Carpathian (Alpine) thrust front, RS = Renohercynian suture, STS = Saxothuringian suture, SMS = Staré Město suture (= Moldanubian thrust). (B) Variscan orogen architecture proposed by Franke & Zelazniwicz (2002): EFZ = Elbe fault zone, ISFZ = Intra-Sudetic fault zone.

Similarly, Franke & Zelazniwicz (2000, 2002) emphasize the significant role of the dextral transpression, without the intention of interpreting subduction zones. Unlike the other interpretations, their view is more descriptive (Fig. 3-4B), identifying, for example, at least two Permo-Carboniferous pull-apart basins in Poland. Note that these interpretations, which place a main suture zone well north of the Erzgebirge need not be incompatible with the microplate tectonics discussed by Kröner et al. (2007) and others.

While pre-Variscan and Variscan subduction zones differs in all models mentioned above, and none of them fits all the observed field data, all authors agree on the importance of post-Variscan (i.e., post-collisional) dextral strike-slip tectonics. The onset of this dextral shearing differs between the models however, which potentially has important implications for the mineral systems in the eastern Erzgebirge.

Significantly, all of these models focus mainly on explaining the accretionary phase of the Variscan orogen; while none sufficiently explain the post-peak exhumation, significant dextral strike-slip and (potential) orocline formation. In the context of this project, it is these late to post-variscan processes that are most relevant, and will be investigated in more detail. That said, the subduction zone architecture and accretionary processes that formed the Bohemian massif need to be considered by any mineral systems model, as these likely set up the architecture, fluid-pathways and potentially metallogenic fertility of the broader Erzgebirge region. Other post-orogenic processes, such as delamination or slab-breakoff, would also have significant implications for the Erzgebirge mineral system, though are generally discussed only in vague terms in the existing literature.

Finally, we note that all of the tectonic interpretations described above are based largely on geological mapping and other surface observations, resulting in significant uncertainty beneath recent sedimentary basins (e.g., the central European basin). While various geophysical data sets are available to provide key additional constraints on the crustal (and lithospheric) architecture of the Bohemian massif, we suggest that this has largely been under-utilized.

### *3.1.2. Data compilation*

A variety of published or otherwise publicly available geological, geophysical, geochronological and mineral occurrence data have been compiled into a GIS data package as a basis for synthesis and interpretation. This has been complemented by preliminary geophysical results from the LfULG, which are proving very useful. Key datasets that are included:

- Mapped surface geology from Czech and German agencies
- Gravity + magnetic data
- Crustal seismic-reflection surveys
- Regional scale seismic tomography
- Mineralized vein occurrences and orientations
- Radiometric sampling sites with ages
- Interpreted and mapped faults
- Gamma ray spectrometry maps
- Georeferenced figures from publications (e.g., the figures in Section 3.1.1, , published ages, etc.)

We are in the process of applying a variety of standard processing techniques to the available geophysical data (e.g., upward continuations, vertical derivatives, tilt derivatives) to enhance potential subtle features. Of key interest will be the signatures of potentially ore-related structures, including (1) crustal scale boundaries such as the Elbe shear zone, pull-apart basins and tectonic terranes, and (2) local-scale structures such as intrusions, metamorphic aureole's, and faults that potentially relate to ore deposits.

At the conclusion of the project, a series of thematic maps, cross-sections and interpretations derived from this dataset will be prepared for submission to LfULG, visualized in a way that best communicates our main findings.

### 3.1.3. *Synthesis and preliminary interpretations*

So far, our efforts have focused on understanding the post-peak metamorphism Variscan transtensional structures, such as the Elbe Shear Zone (ESZ), and their links to magmatism and associated mineralization. The ESZ is a major geological structure that forms the prominent eastern border of the Erzgebirge, separating the well-endowed Erzgebirge from the Lausitz Block, which contains only scant magmatic-hydrothermal mineralization.

The ESZ is well known to be a long-lived (and still ongoing) tectonic structure, with mostly dextral kinematics and a total (interpreted) offset of ~40–50 km. It has been subject to a complex interplay of late to post-Variscan tectonic, magmatic, and sedimentary processes, which have left a distinct signature in the magmatic and sedimentary rock record. At least two distinct transtensional events can be distinguished and attributed to the dextral activity on the ESZ. These are recorded by (1) the Meissen Massif, a complex intrusion that was emplaced at 330–320 Ma (Sharp et al., 1997; Nasdala et al., 1999) in a dilatational jog (Kröner et al., 2007), and (2) the Döhlen Basin near Freital, the volcano-sedimentary infill of a small pull-apart basin with abundant ignimbrite/tuff units ranging in age from ~294–286 Ma. The latter evidence for a tectonic reactivation along the ESZ and synchronous magmatic events (Zieger et al., 2019).

These ages bracket a potentially extensive period of activity of the ESZ, supporting the potential for a spatio-temporal link between the ESZ and formation of the Tharandt and Altenberg-Teplice Calderas at ~315–310 Ma (Tomek et al., 2021). The latter is well known to host a number of small granitic stocks with related Sn-Li/W-(Mo) mineralization. Significantly, these zones of strain-transfer between NW–SE trending dextral strike-slip faults localize felsic magmatism, by accommodating intrusions and establishing favorable pathways to the lower crust through which melts can ascend. This interpretation is supported by the generally A-type granite composition of mineralized intrusions (Klomínský et al., 2010), which commonly occur along shear zones during within-plate transtension (e.g., Collins et al., 2019).

Recognized epithermal veins and greisen-type Sn, and W mineralization in the Eastern Erzgebirge are mostly of Paleozoic age and, with better geochronology, can likely be linked with specific tectonic and/or magmatic events. Precise dating of intra- and extra- ATC caldera rhyolites and ring dykes (Tomek et al., 2021; Tichomirowa et al., 2022; Casas-Garcia et al., 2019) allows direct linking to Sn deposits of the same age (Seifert et al., 2011; Leopardi et al., *submitted*). The role of contemporaneous dextral wrenching tectonics in the Bohemian Massif (Edel et al., 2021; Bárta et al., 2021) remains unclear, but we suggest likely structural controls the localization of pluton emplacement and their accommodation and growth. An analogous



situation can be observed in the western Erzgebirge (albeit likely at a deeper crustal level), where granitic plutonism and associated Sn and W mineralization is clearly associated with major NW trending dextral strike-slip structures (Fig. 3-5).

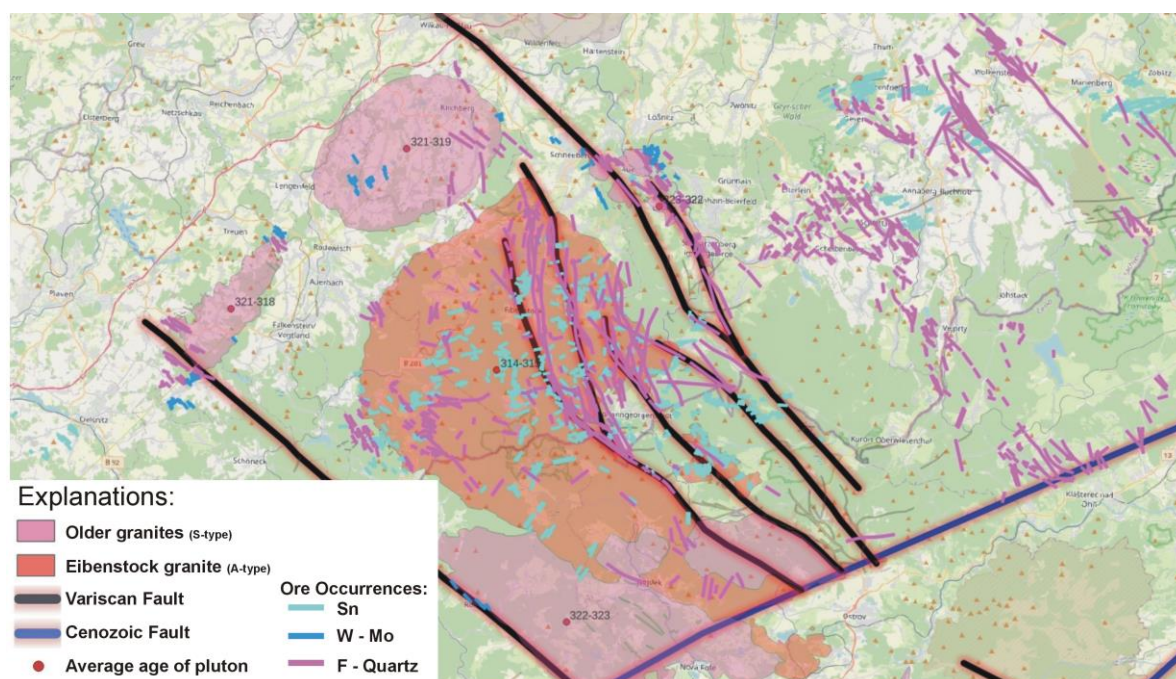


Figure 3-5. Tectonic situation on the eastern margin of the Eibenstock pluton and related ore deposits. Plutons clearly sit within a transtensional jog (and possible flower structure) within the broader Gera-Jáchymov fault system. Significantly, this fault system appears to also control the major overprinting mineralized F-Quartz vein systems.

It is likely that the same tectonic regime applied also to the Eastern Erzgebirge, which is supported by previously provided references and summarized in the tectonostratigraphic chart for the Paleozoic Era for the Erzgebirge region (Fig. 3-6). The similar age of the Eibenstock pluton and Altenberg-Teplice caldera (ATC) strongly suggests that magmatic system beneath the caldera was emplaced in the same setting, and potentially that similar fault systems can be expected at depth. Comparison of the Western and Eastern Erzgebirge (and considering their different erosional level) reveals two additional potential analogies: (1) one might expect significant fluorite vein deposits along emplacement-related faults beneath the ATC, as in the Western Erzgebirge, and; (2) W-Mo mineralization could be associated with relatively older intrusions crosscut by Sn-rich ones at ~315 Ma, as is observed at Eibenstock (Fig. 3-5 and 3-7).



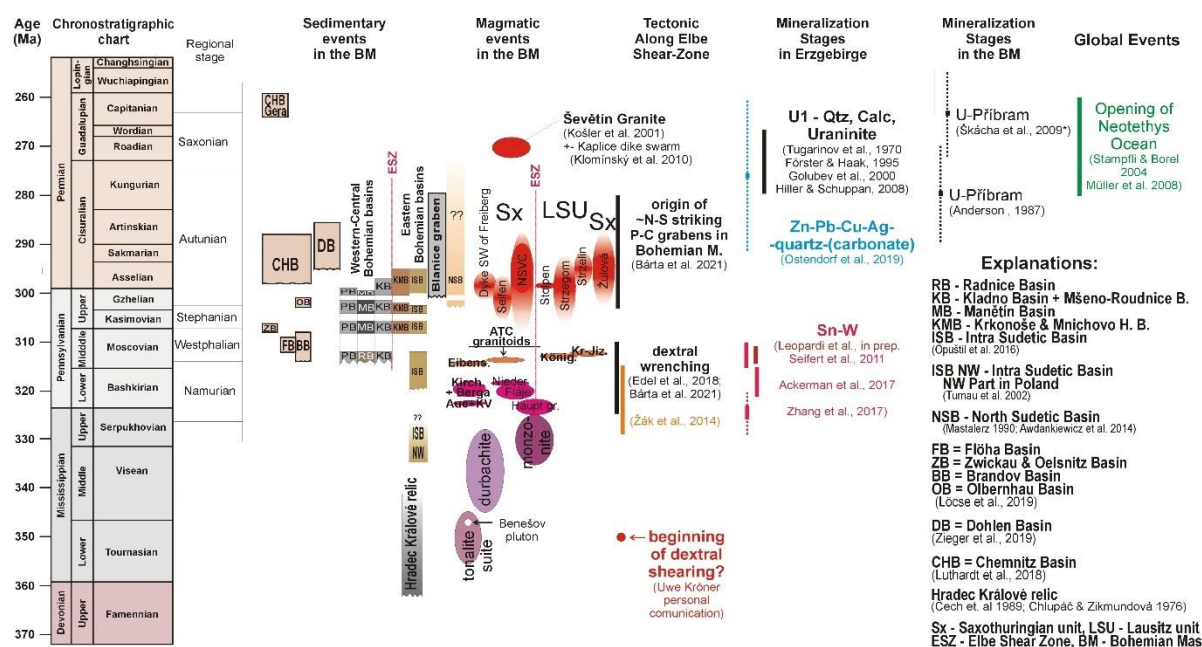


Figure 3-6. Tectonostratigraphic chart with ages of ore deposits for the Paleozoic Era in the BM. The compiled chart is based on following references: Tectonic events – Bárta et al. (2021), Edel et al. (2018), Žák et al. (2014); Magmatic events – Bárta et al. (2021), Breitkreuz et al. (2021), Casas-García et al. (2019) Förster & Rhede (2006), Hoffmann et al. (2009, 2013), Lócse et al. (2019), Kässner et al. (2021), Klomínský et al. (2010), Košler et al. (2001), Kryza et al. (2014), Laurent et al. (2014), Nasdala et al. (1999), Oberc-Dziedzic & Kryza (2012), Sharp et al. (1997), Tichomirowa et al. (2019, 2022), Tomek et al. (2021), Turniak et al. (2014), Wenzel et al. (1997); Mineralization events – Ackerman et al. (2017), Anderson (1987), Förster & Haak (1995), Golubev et al. (2000), Hiller & Schuppan (2008), Ostendorf et al. (2019), Seifert et al. (2011), Škácha et al. (2009), Tugarinov et al. (1970), Zhang et al. (2017); Sedimentary events – Awdankiewicz et al. (2014), Chlupáč & Zikmundová (1976), Čech et al. (1989), Lócse et al. (2019), Luthardt et al. (2018), Mastalerz (1990), Opuštil et al. (2016), Turnau et al. (2002), Zieger et al. (2019); Global events – Stampfli & Borel (2004), Müller et al. (2008).

Known W-Mo mineralization around relatively older Telnice granite and northern part of the Fláje pluton in the Eastern Erzgebirge have been reported (Uwe Lehmann, personal communication), while Sn mineralization is well known around intrusions (312-314 Ma, and possibly ~301 ±5 Ma near Seiffen; Th-U-Pb on monazite & K-Ar on Siderophyllite, Förster & Rhede 2006). Fluorite-quartz veins are also reported from Sadisdorf (Dino Leopardi, personal communication), although their timing is unclear. Lastly, our analogy to the Western Erzgebirge would imply that published ages for the Markersbach (ICP-MS; 327 ±4 Ma; Hofmann et al., 2009) and most likely also the Telnice granite (K-Ar; 338 Ma; Štemprok et al., 2003) are likely incorrect, because the Markersbach granite penetrates the Elbe shear zone and, therefore, it should be one of the youngest intrusions in the area (certainly younger than the Meissen Massif). Moreover, associated Sn mineralization occurrences around Markersbach suggest that this pluton should be of similar age as the ATC granites. The only available radiometric age for the Telnice granite ages is in literature referred to as “personal communication” with no actual data provided (Štemprok et al., 2003). The age is thus deemed unreliable and likely too old, as this would make it the oldest plutonic body in the Eastern

Erzgebirge – essentially coinciding with the age of peak Variscan metamorphism. Furthermore, it would be of the same age as the Central Bohemian Plutonic Complex. Moreover, the fact that W-Mo mineralization is related to the Telnice granite may be used to suggest that this granite is at least 20 Ma younger than the currently available age.

From, it is suggested to obtain new radiometric ages for the Markersbach granite, the Telnice granite – as well as the currently undated Schellerhau granite in order to provide a more consistent magmatic and geotectonic Framework for the Eastern Erzgebirge and its magmatic-hydrothermal mineral systems.

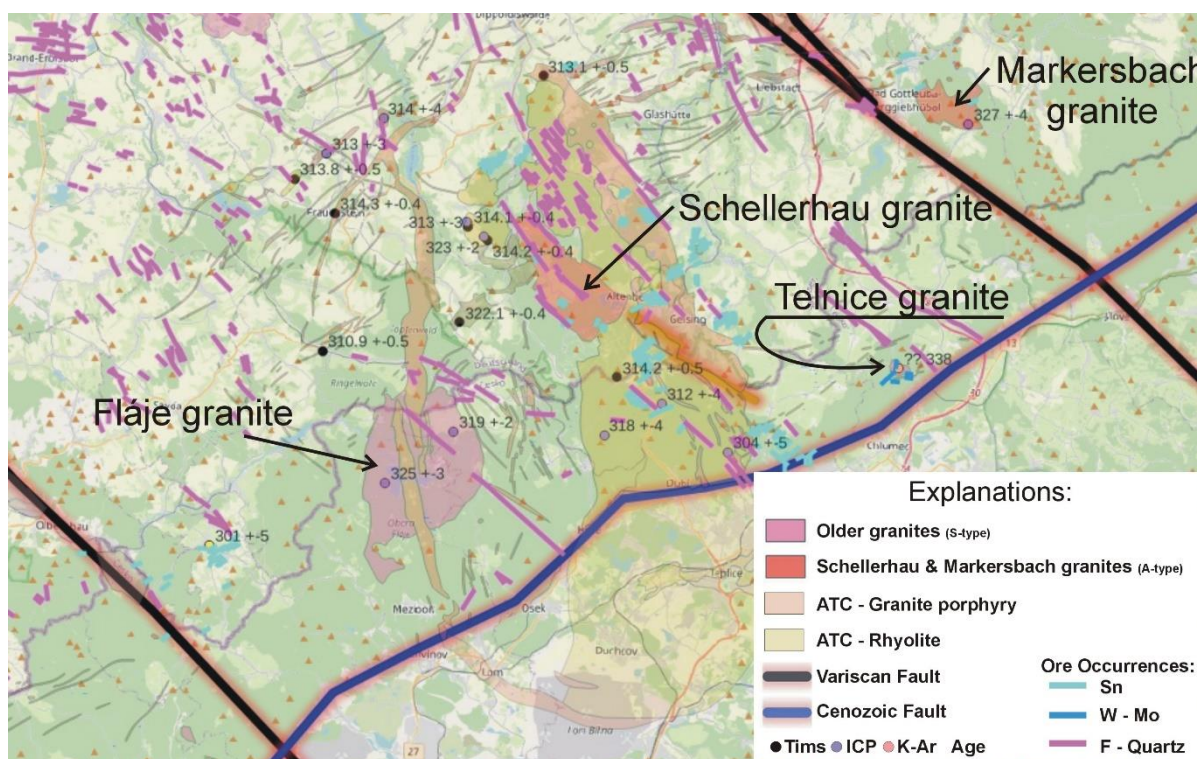


Figure 3-7. Tectonic situation on the Eastern Erzgebirge study region and related magmatic-hydrothermal ore deposits.

### 3.3. Next steps

At the current stage, structural data from historic maps and other publications (Baumann, Hofmann, etc.) are still being collated for the magmatic and epithermal systems in the eastern Erzgebirge. Once complete, these will provide the foundation for a structural analysis of the fluid pathways that controlled mineralization, and potentially the broader stress regime during mineralization.

Newly provided magnetic and gravity data are being analyzed and (re)-processed (e.g., high-pass filtering, gradients) to allow a joint geology and geophysical interpretation of

the eastern Erzgebirge. These will provide a basis to better evaluate published tectonic models for the region and the link between tectonics and mineralization.

## 4. WP 3 to 5 - Mineral Systems Analysis

These work packages will be the focus of future reports, once more progress has been made in WP-2 and WP-3.

However, to help place our preliminary work in context, we have included a brief conceptual model for the Eastern Erzgebirge mineral system (Fig. 4-1). This will evolve significantly over the course of this project, but hopefully serves as a useful starting point for discussions. It highlights key elements that likely control the epithermal and magmatic mineral systems, and the (potential) commonalities between them. While a detailed review of the literature on the metallogensis of the Freiberg district is out-of-scope for this progress report, the many question marks in the diagram emphasize that significant uncertainty exists.

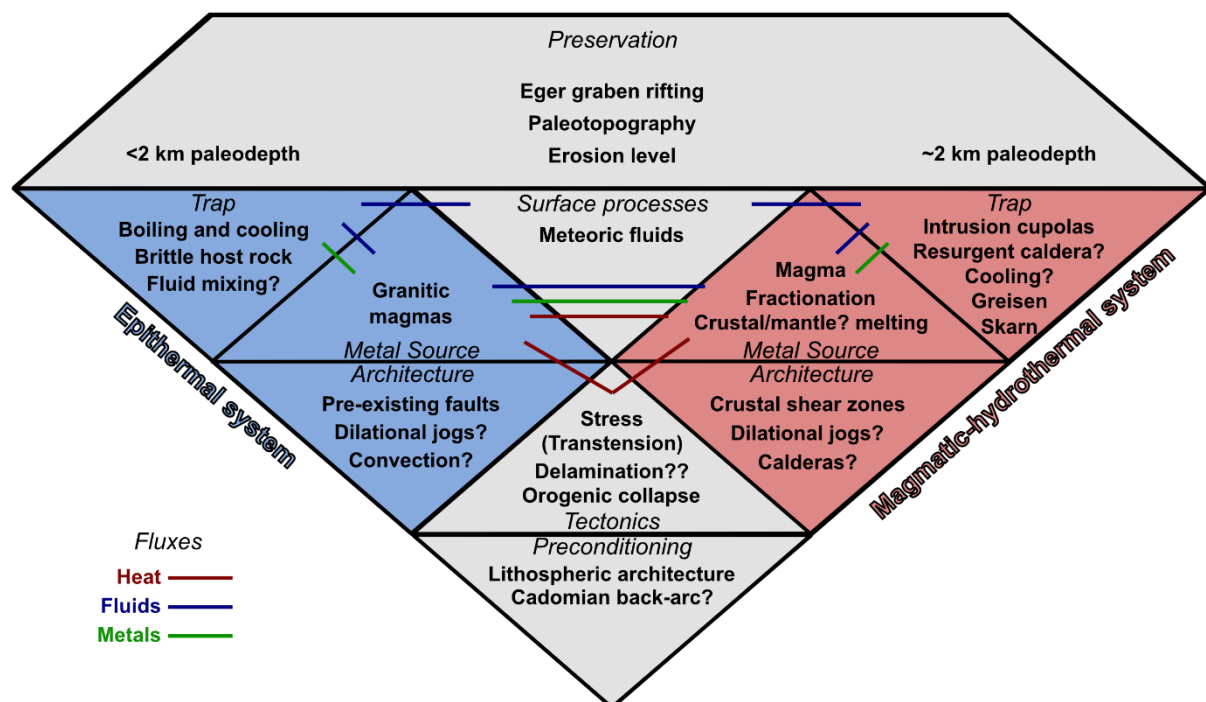


Figure 4-1. Summary diagram highlighting the key components and fluxes that are likely to play a role in the Eastern Erzgebirge mineral systems. Common elements are shown in grey, epithermal-specific processes in blue and magmatic-specific processes in red. The main fluxes are also schematically overlain as colored arrows.

This conceptual model has allowed us to develop several working hypotheses for the Paleozoic mineral systems in the Erzgebirge:

1. Sn-Li-W-Mo deposits can be divided into two subtypes: (1) Sn-Li-dominated and (2) W-Mo-dominated deposits. Both these subtypes are clearly directly linked with magmatic systems. The Sn-Li-dominated deposits occur in greisens, greisenised lenses, veins, and stockworks, while the W-Mo-dominated mineralization is almost solely related to quartz veins and stockworks (e.g., Štemprok & Blecha, 2015). The geographical distribution of Sn-Li and W-Mo dominated deposits differ, suggesting either different mobility, or association with magmatism of different ages (312–315 Ma vs. 318–322 Ma). Nevertheless, the Sadisdorf deposit contains not only Sn and Li, but also Mo. This may be explained by a higher susceptibility of Mo for remobilization due to thermo-chemical sulfate reduction (i.e.,  $80 < T < 200$  °C; Ardakani et al., 2016).
2. Epithermal Pb-Zn-Ag: Although often inferred, a magmatic system that can be directly related to this/these deposits remains unknown. Fluid inclusion analyses has shown that fluids are partly of magmatic origin and partly of hydrothermal origin (Burisch et al., 2019b; Bauer et al., 2019; Swinkels et al., 2021), as is typical in epithermal systems. Various factors likely control the emplacement of these deposits, including:
  1. A suitable host rock to facilitate brittle vein formation
  2. Pre-existing structures oriented favorably for fluid flow relative to the stress field (e.g., dilational jogs)
  3. Vicinity of a suitable parental magmatic system. Magmatic rocks possibly related to such magmatic system and their ages are listed in the supplementary table (Appendix 1)
  4. Suitable source rocks which might have fertilized magma
  5. Depth below the paleosurface and preservation potential due to subsequent uplift and erosion

So far, it is unclear which of these factors played the most significant role in the eastern Erzgebirge, nor how they can be used to predict undiscovered mineralization. Nevertheless, we can speculate that the mineralized veins originated during (1) trans-tensional or (2) extensional regimes following the collapse or dismemberment of the Variscan orogen. Further analyses will be focused on resolving these essential questions.



## 5. WP 6 - Publication and dissemination

Two abstracts covering the preliminary results of work packages 1 and 2 have been submitted for the 10<sup>th</sup> Hutton Symposium on granites and related rocks (10-15 September 2023 in Italy; <https://hutton10.eu/>). As we are in the early stage of the “Neues Potential” project, nothing has been decided yet concerning publication in peer-reviewed international journals, however this is expected over the next ~1 year.

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## 7. Appendix 1

Compilation table of the ages available for magmatic, volcanic, sedimentary and ore deposit events. \*For the confidence level, an age is considered “robust” if the method used is considered reliable in the geochronology community and provide a precise age (often with a rather small analytical error  $\leq 5$  Ma) with a good description of the sample’s mineralogy, geochemistry and location. Ages considered “passable” are determined by less precise method (e.g., Sm-Nd) which often give analytical error  $>5-10$  Ma. A “Questionable” age is in contradiction with other geochronological study or petrographic/fieldwork observations.

Region	Sample	Type	Age	Error	Locality	Sublocality	Method	Remarks	Confidence level*	Reference
<b>Early Permian &amp; Carboniferous basins</b>										
<i>Döhlen Basin</i>										
E. Erzgebirge	Unk2	Unkersdorf Tuff (Unkersdorf Formation)	294	3	Döhlen Basin	Freital (quarry Osterberg near the Weißeritz river)	U–Th–Pb LA-SF ICP-MS zircon	Concordia age (interpreted as the age of emplacement of the tuff by authors)	Robust	Zieger et al. (2019)
E. Erzgebirge	Unk1	Trachyandesite (Unkersdorf Formation)	293	5	Döhlen Basin	Freital (on top of the Burgwartsberg)	U–Th–Pb LA-SF ICP-MS zircon	Concordia age (interpreted as the cooling age by authors)	Robust	Zieger et al. (2019)
E. Erzgebirge	Ban4	Wachtelberg Ignimbrite (Upper Bannewitz Formation)	286	4	Döhlen Basin	Field between Freital and Cunnersdorf	U–Th–Pb LA-SF ICP-MS zircon	Concordia age	Robust	Zieger et al. (2019)
E. Erzgebirge	DB-1	Rhyodacitic lava (Unkersdorf Formation)	296	3	Döhlen Basin	Active quarry south of Wilsdruff	U-Pb SHRIMP zircon	12 zircons	Robust	Hoffmann et al. (2013)
E. Erzgebirge	DB-2	Rhyolitic Zauckerode Tuff, Niederhäslich	292	13	Döhlen Basin	Temporary outcrop in Freital	Pb/Pb single zircon evaporation age		Passable	Hoffmann et al. (2013)

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<i>Chemnitz basin</i>										
E. & C. Erzgebirge		Zeisigwald Tuff pyroclastics	291	2	Chemnitz basin		U-Pb La-ICP-MS zircon		Robust	Luthardt et al. (2018)
E. & C. Erzgebirge	CB-1	Rhyolitic vitrophyric Planitz Ignimbrite (Planitz Formation)	296,6	3	Chemnitz basin, temporary outcrop in Zwickau	Planitz Formation	U-Pb SHRIMP zircon	14 zircons, Wetherhill graph	Robust	Hoffmann et al. (2013)
E. & C. Erzgebirge		Rhyolite dyke	278	5	St. Egidien and Chemnitz		<sup>206</sup> Pb/ <sup>238</sup> U SHRIMP zircon	interpreted as age of postkinematic Late Variscan volcanism in the Sub-Erzgebirge basin	Robust	Nasdala et al. (1998)
<i>Kladno-Rakovník Basin</i>										
E. Erzgebirge	25a	Tuff	297,16	0,17	Kladno-Rakovník Basin	Líne Formation	U-Pb CA-ID-TIMS zircon		Robust	Opluštil et al. (2016)
E. Erzgebirge	24	Clay horizon	298,97	0,09	Kladno-Rakovník Basin	Klobuky coal	U-Pb CA-ID-TIMS zircon		Robust	Opluštil et al. (2016)
E. Erzgebirge	16	Clay horizon	301,5	0,11	Kladno-Rakovník Basin	Kamínek	U-Pb CA-ID-TIMS zircon		Robust	Opluštil et al. (2016)
E. Erzgebirge	17a	Tuff	302,47	0,08	Kladno-Rakovník Basin	Mšec Member	U-Pb CA-ID-TIMS zircon		Robust	Opluštil et al. (2016)
E. Erzgebirge	11	Tuff	307,05	0,16	Kladno-Rakovník Basin	Nýřany Member	U-Pb CA-ID-TIMS zircon		Robust	Opluštil et al. (2016)
E. Erzgebirge	10	Tuff	312,36	0,07	Kladno-Rakovník Basin	Upper Radnice Member (Black Tuffite)	U-Pb CA-ID-TIMS zircon		Robust	Opluštil et al. (2016)
E. Erzgebirge	9	Tuff	312,38	0,05	Kladno-Rakovník Basin	Green Tuff	U-Pb CA-ID-TIMS zircon		Robust	Opluštil et al. (2016)

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E. Erzgebirge	8	Tuff	313,22	0,04	Kladno-Rakovník Basin	Z-tuff	U-Pb CA-ID-TIMS zircon		Robust	Opluštil et al. (2016)
E. Erzgebirge	6	Tuff	313,2	0,06	Kladno-Rakovník Basin	Middle Volcanic Horizon	U-Pb CA-ID-TIMS zircon		Robust	Opluštil et al. (2016)
E. Erzgebirge	2	Tuff	313,41	0,07	Kladno-Rakovník Basin		U-Pb CA-ID-TIMS zircon		Robust	Opluštil et al. (2016)
E. Erzgebirge	4	Clay horizon	313,83	0,23	Kladno-Rakovník Basin	Velká Opuka	U-Pb CA-ID-TIMS zircon		Robust	Opluštil et al. (2016)
<i>Pilsen Basin</i>										
E. Erzgebirge	14	Tuff	303,73	0,13	Pilsen Basin	Týnec Formation	U-Pb CA-ID-TIMS zircon		Robust	Opluštil et al. (2016)
E. Erzgebirge	13	Clay horizon	305,99	0,07	Pilsen Basin		U-Pb CA-ID-TIMS zircon		Robust	Opluštil et al. (2016)
E. Erzgebirge	12	Clay horizon	308	0,04	Pilsen Basin		U-Pb CA-ID-TIMS zircon		Robust	Opluštil et al. (2016)
<i>Radnice Basin</i>										
E. Erzgebirge	3	Ash bed	314,36	0,09	Radnice Basin		U-Pb CA-ID-TIMS zircon		Robust	Opluštil et al. (2016)
C. Erzgebirge	BOB-1	Rhyolitic ? (Olbernhau Formation)	302,6	2,8	Fallout deposit (rhyolitic?) (Drilling B6/99, 13.15–13.25 m depth)		U-Pb SHRIMP zircon	7 zircons	Robust	Hoffmann et al. (2013)
<b>North Saxon Volcanic Complex</b>										
E. Erzgebirge		Distal tuff	298,3	9,1		Kohren formation/ Flechtingen	Pb/U SHRIMP on zircon		Passable	Hoffmann et al. (2013)
E. Erzgebirge		Ignimbrite	294,4	1,8		Rochlitz ignimbrite	Pb/U SHRIMP on zircon		Robust	Hoffmann et al. (2013)
E. Erzgebirge		Ignimbrite	285,6	1,1		Wurzen Volcanic System	U-Pb on zircon		Robust	Wendt et al. (1995)

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E. Erzgebirge		Ignimbrite	284	11		Wurzen Volcanic System	Sm/Nd whole rock	Passable	Wendt et al. (1995)
E. Erzgebirge		Ignimbrite	287	5		Wurzen Volcanic System	Rb/Sr	Robust	Wendt et al. (1995)
E. Erzgebirge		Rhyolite	289,3	4,1		Wurzen Volcanic System	Pb/U SHRIMP on zircon	Robust	Hoffmann et al. (2013)
E. Erzgebirge		Rhyolite	290,2	4,1		Wurzen Volcanic System	Pb/U SHRIMP on zircon	Robust	Hoffmann et al. (2013)
<b>Freiberg district dykes</b>									
E. Erzgebirge		Rhyolite	~300	n.d.		Rhyolite dyke	U-Pb CA-ID-TIMS on zircon		Tichomirowa et al. (personal communication)
<b>Epithermal vein</b>									
E. Erzgebirge		Ag-Pb-Zn epithermal vein	276	16	Freiberg		Rb-Sr isotopes sphalerite	Questionable	Ostendorf et al. (2019)
<b>Skarn</b>									
E. Erzgebirge		Skarn	304,9	3,2	Berggießhübel		U-Pb La-ICP-MS garnet	Robust	Burisch et al. (2019a)
E. Erzgebirge		Skarn	301,5	5,7	Berggießhübel		U-Pb La-ICP-MS garnet	Robust	Burisch et al. (2019a)
<b>Younger granite</b>									
E. Erzgebirge		Granite	302	4	Seiffen		Th-U-Pb monazite	Robust	Förster and Rhede (2006)
E. Erzgebirge		Granite	301	5	Seiffen		K-Ar siderite	Robust	Förster and Rhede (2006)
<b>Oederan Forest Volcano</b>									
E. Erzgebirge	MfNC-2014-01	Ignimbrite	309,4	2,6	Oederan Forest Volcano		U-Pb LA-ICP-MS zircon	Robust	Löcse et al. (2019)
E. Erzgebirge	MfNC-2014-02	Ignimbrite	309	1,8	Oederan Forest Volcano		U-Pb LA-ICP-MS zircon	Robust	Löcse et al. (2019)
E. Erzgebirge	MfNC-2014-05	Schweddey ignimbrite	310,4	1,6	Schweddey		U-Pb LA-ICP-MS zircon	Robust	Löcse et al. (2019)

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E. Erzgebirge	MfNC-2014-08	Schweddey ignimbrite	310,1	2	Schweddey	U-Pb LA-ICP-MS zircon	Robust	Löcse et al. (2019)
E. Erzgebirge	MfNC-2014-11	Roter Stein dyke	311,4	2,7	Roter Stein	U-Pb LA-ICP-MS zircon	Robust	Löcse et al. (2019)
E. Erzgebirge	MfNC-2014-12	Metzdorf dyke	298,6	2,5	Metzdorf	U-Pb LA-ICP-MS zircon	Robust	Löcse et al. (2019)

**Flöha basin**

E. Erzgebirge		Flöha ignimbrite	311,8	9,4	Flöha ignimbrite	U-Pb LA-ICP-MS zircon	Robust/passable	Löcse et al. (2019)
E. Erzgebirge		Flöha ignimbrite	309,2	3,8	Flöha ignimbrite	U-Pb LA-ICP-MS zircon	Robust	Löcse et al. (2019)

**Greisen**

*Altenberg*

E. Erzgebirge	ALT-1	Greisen(?)	321,4	3,3	Altenberg	U-Pb LA-ICP-MS cassiterite	Questionable	Zhang et al. (2017)	
E. Erzgebirge	ALT-2	Greisen(?)	322,4	2,8	Altenberg	U-Pb LA-ICP-MS cassiterite	Questionable	Zhang et al. (2017)	
E. Erzgebirge	5000	Greisen (metamonzogranite greisenized)	323,9	2,5	Altenberg	Re-Os molybdenite	Questionable	Romer et al. (2007)	
E. Erzgebirge	TS 1	Quartz vein	317,9	2,4	Altenberg	Re-Os molybdenite	Robust	Romer et al. (2007)	
E. Erzgebirge		Granite?	318	2	Altenberg	Re-Os molybdenite	Questionable	Romer et al. (2007)	
E. Erzgebirge		Granite?	324	2	Altenberg	Re-Os molybdenite	Questionable	Romer et al. (2007)	
E. Erzgebirge		Granite	327	4	Markersbach	U-Pb LA-ICP-MS monazite, xenotime, uraninite	Questionable	Hofmann et al. (2009)	
E. Erzgebirge	Gal 20-39	Granite	305	3	Altenberg	Rb-Sr whole-rock	lowered isochron age	Questionable	Gerstenberger (1989)
E. Erzgebirge		Granite	293	n.d.	Altenberg	Rb-Sr whole-rock	Error not available	Robust	Haack (1990)

*Krupka*



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E. Erzgebirge	KP-1	Greisen	320,1	2,8	Krupka	U-Pb LA-ICP-MS cassiterite	Questionable	Zhang et al. (2017)	
E. Erzgebirge	greisen	Greisen	312	3,1	Krupka-Knötel	<sup>40</sup> Ar/ <sup>39</sup> Ar Li-mica (protolithionite)	Robust	Seifert and Pavlova (2016)	
E. Erzgebirge	51/M	Quartz vein within greisen (metamonzogranite greisenized)	319,2	2	Krupka-Knötel	Re-Os molybdenite	Robust	Ackerman et al. (2017)	
E. Erzgebirge	51/M	Quartz vein within greisen (metamonzogranite greisenized)	317,7	2	Krupka-Knötel	Re-Os molybdenite	Robust	Ackerman et al. (2017)	
E. Erzgebirge	Kru-2	Quartz vein within greisen (metamonzogranite greisenized)	315,3	2,3	Krupka-Preisselberg	Re-Os molybdenite	Robust	Ackerman et al. (2017)	
<i>Zinnwald/Cínovec</i>									
E. Erzgebirge		Massive greisen, greisen veins, non greisenized	312,8	1,8	Zinnwald/Cínovec	<sup>40</sup> Ar/ <sup>39</sup> Ar Li-mica (zinnwaldite)	average inverse isochron age; interpreted by authors as a minimum age for greisenisation	Robust	Neßler et al. (2016)
E. Erzgebirge	Cin-9	Quartz vein within greisen (meta-albite granite greisenized)	322,4	5,5	Zinnwald	Re-Os molybdenite	Questionable	Ackerman et al. (2017)	
E. Erzgebirge	ZW-1	Greisen	321,6	3,1	Zinnwald	U-Pb LA-ICP-MS cassiterite	Questionable	Zhang et al. (2017)	
E. Erzgebirge	N.A.	Greisen	312,6	2,1	Zinnwald	Ar-Ar Li-mica	Robust	Seifert et al. (2011)	
E. Erzgebirge	N.A.	Greisen	314,9	2,3	Zinnwald	Ar-Ar Li-mica	Robust	Seifert et al. (2011)	
E. Erzgebirge		Greisen (meta albite granite greisenized)	314,1	1,5	Zinnwald	Ar-Ar zinnwaldite	Robust	Atanasova (2012)	

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E. Erzgebirge		Granite	309	3	Zinnwald	protholitionite granite	K-Ar in Li-mica	recalculated age no errors given	Robust	Dolejš and Štemprok (2001)
E. Erzgebirge		Greisen	311,4	3	Zinnwald		<sup>40</sup> Ar/ <sup>39</sup> Ar Li-mica (zinnwaldite)	Li-Sn-W deposit	Robust	Seifert and Pavlova (2016)
E. Erzgebirge		Greisen (vein)	312,5	3,1	Zinnwald		<sup>40</sup> Ar/ <sup>39</sup> Ar Li-mica (zinnwaldite)		Robust	Seifert and Pavlova (2016)
E. Erzgebirge		Greisen (albite granite)	321	22	Zinnwald	Cinovec+Sachsen höhe	Sm-Nd fluorite		Passable	Höhndorf et al. (1994)
<i>Sadisdorf</i>										
E. Erzgebirge	Sad-6	Greisen (meta albite granite greisenized)	321,4	3,8	Sadisdorf		Re-Os molybdenite		Questionable	Ackerman et al. (2017)
E. Erzgebirge	SD-1	Greisen/vein	326,1	3,4	Sadisdorf		U-Pb LA-ICP-MS cassiterite		Questionable	Zhang et al. (2017)
E. Erzgebirge		Mica-quartz-topaz greisen ("Innengreisen"), (meta albite granite greisenized)	310	3,5	Sadisdorf	Sn(-W-Mo-Cu-Zn-In) deposit	<sup>40</sup> Ar/ <sup>39</sup> Ar Li-mica (protolithionite)		Robust	Seifert and Pavlova (2016)
E. Erzgebirge		Meta albite granite greisenized	326	8	Sadisdorf		Sm/Nd wolframite		Questionable	Kempe and Belyatsky (1997)
<b>Lamprophyres</b>										
C. Erzgebirge		Minette dike (Lp)	316	3	Griesbach (SW Zschopau)		<sup>40</sup> Ar/ <sup>39</sup> Ar biotite		Robust	von Seckendorff et al. (2004)
C. Erzgebirge		Kersantite dike (Lp)	319	3	Scharfenstein (SE Chemnitz)		<sup>40</sup> Ar/ <sup>39</sup> Ar biotite		Robust	von Seckendorff et al. (2004)
E. Erzgebirge		Minette dike (Lp)	320	3	St. Michaelis (Brand-Erbisdorf)		<sup>40</sup> Ar/ <sup>39</sup> Ar biotite		Robust	von Seckendorff et al. (2004)
E. Erzgebirge		Kersantite dike (Lp)	324	3	Rabenau (S Dresden)		<sup>40</sup> Ar/ <sup>39</sup> Ar biotite		Robust	von Seckendorff et al. (2004)

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E. Erzgebirge		Meta albite granite greisenized	313,04	0,67	Sadisdorf	U-Pb LA-ICP-MS cassiterite	Robust	Leopardi et al. (submitted)
<b>Altenberg-Teplice caldera</b>								
<i>Altenberg ring dikes</i>								
E. Erzgebirge	FT13	Microgranite	312	3	Altenberg-Frauenstein microgranite	U-Pb LA-ICP-MS zircon	Robust	Tomek et al. (2019)
E. Erzgebirge	FT63	Microgranite	312	4	Altenberg-Frauenstein microgranite	U-Pb LA-ICP-MS zircon	Robust	Tomek et al. (2019)
E. Erzgebirge	FT153	Microgranite	312	3	Altenberg-Frauenstein microgranite	U-Pb LA-ICP-MS zircon	Robust	Tomek et al. (2019)
E. Erzgebirge	19Elend	Microgranite	313,1	0,5	Altenberg-Frauenstein microgranite	U-Pb CA-ID-TIMS zircon	Robust	Tichomirowa et al. (2022)
E. Erzgebirge	Frau	Microgranite	314,3	0,4	Altenberg-Frauenstein microgranite	U-Pb CA-ID-TIMS zircon	Robust	Tichomirowa et al. (2022)
E. Erzgebirge		Microgranite	319,2	2,4	Altenberg-Frauenstein microgranite	U/Pb evaporation zircon	Robust	Romer et al. (2010)
E. Erzgebirge	GP	Microgranite	297	13	Altenberg-Frauenstein microgranite	K-Ar amphibole	Passable	Müller et al. (2005)

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E. Erzgebirge	Microgranite	307-309	n.d.	Altenberg-Frauenstein microgranite		$^{40}\text{Ar}/^{39}\text{Ar}$ biotite, hornblende		Robust	Seltmann and Schilka (1995)	
	Microgranite	300-290	n.d.	Altenberg-Frauenstein microgranite		$^{40}\text{Ar}/^{39}\text{Ar}$ K-feldspar		Questionable	Seltmann and Schilka (1995)	
E. Erzgebirge	Microgranite (granite porphyry)	290	5	Altenberg-Frauenstein microgranite		$^{207}\text{Pb}/^{206}\text{Pb}$ evaporation zircon	Interpreted by authors as the age of greisenization and Sn mineralization	Questionable	Kempe et al. (1999)	
E. Erzgebirge	Microgranite (granite porphyry)	333	8	Altenberg-Frauenstein microgranite		$^{207}\text{Pb}/^{206}\text{Pb}$ evaporation zircon	Interpreted by authors as the age of the Altenberg microgranite emplacement	Questionable	Kempe et al. (1999)	
<i>Sayda-Berggießhübel Dike Swarm</i>										
E. Erzgebirge	19Biene	Rhyolitic dike	310,9	0,5	Sayda-Berggießhübel dyke swarm	3. phase	U-Pb CA-ID-TIMS zircon	Robust	Tichomirowa et al. (2022)	
E. Erzgebirge	19Turm	Rhyolitic dike	313,8	4,9	Sayda-Berggießhübel dyke swarm	Frauenstein?	U-Pb CA-ID-TIMS zircon	Robust	Tichomirowa et al. (2022)	
E. Erzgebirge	FT168	Rhyolitic dike	314	4	Sayda-Berggießhübel dyke swarm		U-Pb LA-ICP-MS zircon	concordia age	Robust	Tomek et al. (2021)
E. Erzgebirge	FT170	Pyroclastic dike	313	3	Sayda-Berggießhübel dyke swarm		U-Pb LA-ICP-MS zircon	concordia age	Robust	Tomek et al. (2021)
E. Erzgebirge	FT168	Rhyolite	314	4	Sayda-Berggießhübel		U-Pb LA-ICP-MS zircon	concordia age	Robust	Tomek et al. (2021)

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übel dyke  
swarm

<i>Teplice ignimbrite</i>										
E. Erzgebirge	TAVC-2	Igimbrite	308,8	4,9	Teplice ignimbrite		U-Pb SHRIMP zircon	Wetherhill graph	Passable	Hoffmann et al. (2013)
E. Erzgebirge	Sö30	Igimbrite	314,2	0,4	Teplice ignimbrite	Teichweg member	U-Pb CA-ID-TIMS zircon		Robust	Tichomirowa et al. (2022)
E. Erzgebirge	19Hart	Igimbrite	314,1	0,4	Teplice ignimbrite	Lugstein-Pramenáč	U-Pb CA-ID-TIMS zircon		Robust	Tichomirowa et al. (2022)
E. Erzgebirge	19Lug	Igimbrite	314,2	0,5	Teplice ignimbrite	Lugstein-Pramenáč	U-Pb CA-ID-TIMS zircon		Robust	Tichomirowa et al. (2022)
E. Erzgebirge	2055-85	Igimbrite	323	2	Teplice ignimbrite	Teichweg member	U-Pb LA-ICP-MS zircon	concordia age	Questionable	Casas-García et al. (2019)
E. Erzgebirge	HS-01	Igimbrite	313	3	Teplice ignimbrite	Teichweg member	U-Pb LA-ICP-MS zircon	concordia age	Robust	Casas-García et al. (2019)
E. Erzgebirge	TR-22	Igimbrite	310	4	Teplice ignimbrite	Lugstein-Pramenáč	U-Pb LA-ICP-MS zircon	concordia age	Questionable	Casas-García et al. (2019)
E. Erzgebirge	TR-39	Igimbrite	313	5	Teplice ignimbrite	VIčí kámen-Medvědí vrch Member.	U-Pb LA-ICP-MS zircon	concordia age core	Robust	Casas-García et al. (2019)
E. Erzgebirge	TR-39	Igimbrite	304	5	Teplice ignimbrite	Medvědí vrch	U-Pb LA-ICP-MS zircon	concordia age rim	Questionable	Casas-García et al. (2019)
E. Erzgebirge	TR-19	Igimbrite	312	4	Teplice ignimbrite	Přední Cínovec member	U-Pb LA-ICP-MS zircon	concordia age	Robust	Casas-García et al. (2019)
E. Erzgebirge	FT164	Igimbrite	314	3	Teplice ignimbrite	extra caldera facies	U-Pb LA-ICP-MS zircon	concordia age	Robust	Tomek et al. (2021)
E. Erzgebirge	FT164	Igimbrite	314	3	Teplice ignimbrite	Přední Cínovec member	U-Pb LA-ICP-MS zircon	concordia age	Robust	Tomek et al. (2021)
<i>Schönfeld complex</i>										
E. Erzgebirge	19Schön	Igimbrite	322,1	0,4	Schönfeld		U-Pb CA-ID-TIMS zircon		Robust	Tichomirowa et al. (2022)
E. Erzgebirge	SBR	Rhyolite	300	11	Schönfeld	Schönfeld basal rhyolite	K-Ar biotite		Passable	Müller et al. (2005)
E. Erzgebirge	TAVC-1	Rhyolitic Igimbrite	326,8	4,3	Schönfeld	Mikulov Beds	U-Pb SHRIMP zircon	Wetherhill graph	Robust	Hoffmann et al. (2013)

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Tharandt Forest Caldera										
E. Erzgebirge	Grund	Ignimbrite	311,9	0,4	Tharand Forest Caldera		U-Pb CA-ID-TIMS zircon		Robust	Breitkreuz et al. (2021)
E. Erzgebirge	99-103	Ignimbrite	313,4	0,4	Tharand Forest Caldera		U-Pb CA-ID-TIMS zircon		Robust	Breitkreuz et al. (2021)
E. Erzgebirge	99-63	Rhyolite	314,5	0,5	Tharand Forest Caldera	Outer ring dike	U-Pb CA-ID-TIMS zircon		Robust	Breitkreuz et al. (2021)
Meißen volcanic complex										
E. Erzgebirge	MVC-1	Ignimbrite	302,9	2,5	Meißen massif	Rhyodacitic Leutewitz Ignimbrite	U-Pb SHRIMP zircon	2 zircons	Passable	Hoffmann et al. (2013)
Pre-caldera Plutons										
E. Erzgebirge	Nbg 1	Granite	318,2	0,5	Niedebobri sch		U-Pb CA-ID-TIMS zircon		Robust	Breitkreuz et al (2021)
E. Erzgebirge	Nbg 2	Granite	319,5	0,4	Niedebobri sch		U-Pb CA-ID-TIMS zircon		Robust	Breitkreuz et al (2021)
E. Erzgebirge		Granite	324	4	Niedebobri sch		U-Pb dating by isotope dilution-thermal ion mass spectrometry monazite, xenotime, uraninite		Passable	Förster et al. (1998)
E. Erzgebirge		Granite	320	6	Niedebobri sch		Pb/Pb evaporation zircon		Passable	Tichomirowa (1997)
E. Erzgebirge	FLG01	Monzogranite	319	2	Fláje		U-Pb LA-ICP-MS zircon	concordia age	Robust	Tomek et al. (2021)
E. Erzgebirge	FT163	Monzogranite	326,8	4,3	Fláje		U-Pb LA-ICP-MS zircon	concordia age	Robust	Tomek et al. (2021)
E. Erzgebirge	FLG01	Monzogranite	319	2	Fláje		U-Pb LA-ICP-MS zircon	concordia age	Robust	Tomek et al. (2021)
E. Erzgebirge	FT163	Monzogranite	325	3	Fláje		U-Pb LA-ICP-MS zircon	concordia age	Robust	Tomek et al. (2021)
E. Erzgebirge		Pluton	326	6	Leuben		U-Pb SHRIMP zircon		Passable	Nasdala et al. (1999)



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E. Erzgebirge		Pluton	330	5	Heidensch anze		U-Pb SHRIMP zircon	Passable	Nasdala et al. (1999)
E. Erzgebirge		Pluton	330,4	2,8	Leuben		Ar-Ar amphibole	Robust	Wenzel et al. (1997)
E. Erzgebirge		Pluton	329,1	2,8	Heidensch anze		Ar-Ar amphibole	Robust	Wenzel et al. (1997)
E. Erzgebirge		Pluton	323,4	1	Hauptgrani te		?	Robust	Sharp et al. (1997) in Müller (2011)
E. Erzgebirge		Pluton	323,6	1	Riesenstei ngranit		?	Robust	Sharp et al. (1997) in Müller (2011)
<b>Central Erzgebirge Plutons</b>									
C. Erzgebirge		Greifenstein granite	323,9	3,5	Ehrenfried ersdorf		<sup>206</sup> Pb/ <sup>238</sup> U ID TIMS (dating by isotope dilution-thermal ion mass spectrometry) uraninite	Robust	Romer et al. (2007)
C. Erzgebirge	Q8	Sauberg stockscheider (marginal flat-lying pegmatite) aplite	320,6	1,9	Ehrenfried ersdorf	Sauberg mine	<sup>206</sup> Pb/ <sup>238</sup> U ID TIMS uraninite (representing a melt pocket)	Robust	Romer et al. (2007)
C. Erzgebirge		Sauberg stockscheider (marginal flat-lying pegmatite) aplite	319,7	3,4	Ehrenfried ersdorf	Sauberg mine	Th-U-Pb uraninite	Robust	Romer et al. (2007)
C. Erzgebirge		Granite	323,9	2,9	Ehrenfried ersdorf		U-Pb hydrothermal apatite	Robust	Romer et al. (2007)
C. Erzgebirge	E1 + E2	Granite	317,6	1,6	Ehrenfried ersdorf	Sauberg mine	U-Pb hydrothermal apatite	Robust	Romer et al. (2007)
C. Erzgebirge		Ehrenfriedersdor f granite	322	5	Ehrenfried ersdorf		Pb/Pb zircon	Robust	Seifert (2008)
C. Erzgebirge		Ehrenfriedersdor f granite	324-317	n.d.	Ehrenfried ersdorf		U/Pb apatite	Robust	Romer et al. (2007)
C. Erzgebirge		Ehrenfriedersdor f granite	310,8	1,1	Ehrenfried ersdorf		Ar/Ar muscovite	Robust	Werner and Lippolt (1998)
C. Erzgebirge		Pobershau g ranite	321	3	Pobershau		U/Pb ID TIMS monazite	Robust	Warkus et al. (1998)

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C. Erzgebirge	Pobershau granite	317	2	Pobershau	U/Pb monazite, xenotime, uraninite	Robust	Förster (1998)
C. Erzgebirge	Hora Svaté Kateřiny granite	308	14		Th-U-Pb monazite	Passable	Breiter (2008)

West Erzgebirge Plutons									
W. Erzgebirge	ASB 665	Aue granite	324,3	5,5		U-Pb LA-ICP-MS on zircon	<sup>206</sup> Pb/ <sup>238</sup> U weighted mean	Robust	Tichomirowa et al. (2019)
W. Erzgebirge		Aue granite	328,6	2	Aue	Pb/Pb evaporation zircon		Robust	Tichomirowa and Leonhardt (2010)
W. Erzgebirge		Schalenberg granite	325,7	7	Lauter	Pb/Pb evaporation zircon		Robust	Tichomirowa and Leonhardt (2010)
W. Erzgebirge	ASB 668	Lauter granite	326,8	3,6		U-Pb LA-ICP-MS on zircon	<sup>206</sup> Pb/ <sup>238</sup> U weighted mean	Robust	Tichomirowa et al. (2019)
W. Erzgebirge		Lauter granite	331	2,9	Lauter	Pb/Pb evaporation zircon		Robust	Tichomirowa and Leonhardt (2010)
W. Erzgebirge		Abertamy granite	322,8	3,5		Pb/Pb evaporation zircon		Robust	Kovarikova et al. (2007)
W. Erzgebirge	BER 788	Bergen granite	315,8	3,7	Quarry Streuberg	U-Pb LA-ICP-MS on zircon	<sup>206</sup> Pb/ <sup>238</sup> U weighted mean	Robust	Tichomirowa et al. (2019)
W. Erzgebirge	BER 789	Bergen granite	321,1	3,7	Quarry NW Trieb	U-Pb LA-ICP-MS on zircon	<sup>206</sup> Pb/ <sup>238</sup> U weighted mean	Robust	Tichomirowa et al. (2019)
W. Erzgebirge	BER790	Bergen granite	330,3	3,8	Quarry SE Trieb	U-Pb LA-ICP-MS on zircon	<sup>206</sup> Pb/ <sup>238</sup> U weighted mean	Robust	Tichomirowa et al. (2019)
W. Erzgebirge	BER 791	Bergen granite	325,2	4,3	Quarry Kuxenberg	U-Pb LA-ICP-MS on zircon	<sup>206</sup> Pb/ <sup>238</sup> U weighted mean	Robust	Tichomirowa et al. (2019)
W. Erzgebirge		Bergen granite	323	2	Bergen	U/Pb monazite, xenotime, uraninite		Robust	Förster (1998)
W. Erzgebirge		Kirchberg granite	324	5	Kirchberg	U/Pb monazite, xenotime, uraninite		Robust	Förster (1998)
W. Erzgebirge		Kirchberg granite	330	5	Kirchberg	Th-U-Pb uraninite		Robust	Kempe (2003)
W. Erzgebirge		Auerhammer granite	322	6		U-Pb ID TIMS monazite		Passable	Warkus et al. (1998)
W. Erzgebirge	EIB 710	Eibenstock granite	315,1	1,0	Kamelfelsen	<sup>40</sup> Ar- <sup>39</sup> Ar on mica		Robust	Tichomirowa et al. (2019)

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W. Erzgebirge		Eibenstock granite	319,8	1	Eibenstock	Pb/Pb evaporation zircon		Robust	Tichomirowa and Leonhardt (2010)
W. Erzgebirge		Eibenstock granite	316	3	Eibenstock	U-Pb monazite, xenotime, uraninite		Robust	Förster (1998)
W. Erzgebirge		Eibenstock granite	321	4	Eibenstock	U-Pb monazite		Robust	Warkus et al. (1998)
W. Erzgebirge		Eibenstock granite	320	8	Eibenstock	Pb/Pb evaporation zircon		Robust	Kempe et al. (2004)
W. Erzgebirge		Eibenstock granite	322	3	Gottesberg	U/Pb SHRIMP		Robust	Seifert (2008)
W. Erzgebirge		Eibenstock granite	325	3,7	Gottesberg	Pb/Pb		Robust	Seifert (2008)
W. Erzgebirge		Blauenthal granite	315	10		Pb/Pb evaporation zircon		Robust	Tichomirowa and Leonhardt (2010)
W. Erzgebirge		Walfischkopf granite	314,7	4,3		Pb/Pb evaporation zircon		Robust	Tichomirowa and Leonhardt (2010)
W. Erzgebirge		Krinitzberg granite	318,8	2,9		Pb/Pb evaporation zircon		Robust	Tichomirowa and Leonhardt (2010)
W. Erzgebirge		Rhyolite dyke in Gottesberg granite	304-290	n.d.	Eibenstock	U/Pb, K/Ar monazite, biotite		Robust	Förster et al. (2007)
W. Erzgebirge		Rhyolite dyke in Eibenstock	290	5	Eibenstock	Pb/Pb evaporation zircon		Robust	Kempe et al. (2004)
W. Erzgebirge		Rhyolite dyke in Eibenstock	297	8	Eibenstock	U/Pb SHRIMP zircon		Robust	Kempe et al. (2004)
W. Erzgebirge		Microgranite in Eibenstock granite	312,5	4,6	Gottesberg	Pb/Pb evaporation zircon		Robust	Seifert (2008)
W. Erzgebirge	ASB 670	Schwarzenberg granite	322,0	1,4	Schwarzenberg	U-Pb CA-ID-TIMS on zircon	Concordia age	Robust	Tichomirowa et al. (2019)